

A focus in the North Atlantic in the ORCA05-PISCES experiment

R. Dussin, A.M. Treguier, O. Aumont *

LPO Report 2009-09, July 2009

1 Introduction

The global coupled biogeochemistry-dynamics experiment ORCA05-PISCINT3 has shown to have several problems in the North Atlantic. Biogeochemical models like PISCES have a great sensitivity to the dynamics, so that small discrepancies between the model and the real ocean can produce big errors in the biogeochemical variables. The purpose of this report is to document the errors in two fields (nutrients, chlorophyll) and to try to relate them, if possible, to errors in the physical variables (temperature, salinity, density). We will also compare the physics of this ORCA05 simulation (ORCA global grid at 1/2 degree) with simulations at higher resolution carried on by the DRAKKAR project (ORCA025 and NATL12).

ORCA05-PISCINT3 is an experiment forced using ERA40. The total length is 48 years, from 1958 to 2001. The PISCES model is run "online". Initial conditions for tracers are derived from the World Ocean Atlas (Levitus, 2005). For the biogeochemical variables, we have chosen to work with an interannual mean of each month in order to keep the seasonal cycle. The period of interest is 1996-2000, chosen because it is far enough from the beginning of the run to avoid spin-up or adjustments effects and long enough to avoid a singularity of a particular year. All the figures presented here are means of 1996 – 2000 and, for clarity, we chose to focus on some relevant month (march, april, may and august) rather than all the months.

Section 2 is a comparison of surface chlorophyll with data from SEAWIFS, section 3 deals with the mixed layer depth compared to the climatology based on ARGO floats done by Clément de Boyer Montégut. In section 4, we compare the temperature and nitrates with data from the World Ocean Atlas. Finally, section 5 is a comparison with one of the DRAKKAR runs ORCA05-G70.114 and we focus on the differences in the dynamics. Additional figures are provided in 4 annexes. Annex 1 shows results from a Mercator simulation using degraded fields of ORCA025 and PISCES offline. Annex 2 deals with the differences in SST and SSS compared to World Ocean Atlas. Annex 3 focus on the oxygen at 100 meters and along sections. Annex 4 are chlorophyll plots with a linear scale for readers who does not feel comfortable with the logarithmic scale used in section 2. Finally, Annex 5 is the namelist of ORCA05-PISCINT3.

*Laboratoire de Physique des oceans, CNRS-Ifrermer-UBO, Brest, France

2 Comparison of chlorophyll with SEAWIFS

In this section, we compare the mean surface chlorophyll from the ORCA05-PISCINT3 model with a climatology based on Seawifs, computed with data of years 1999 to 2005. Figures 1 and 2 are surface maps of the decimal logarithm of chlorophyll concentration in $mg.m^{-3}$, figure 3 shows surface chlorophyll concentration along zonal sections at $30^{\circ}N$ and $40^{\circ}N$ in the model and in the climatology. Maps of surface chlorophyll concentration with a linear scale are available in annex.

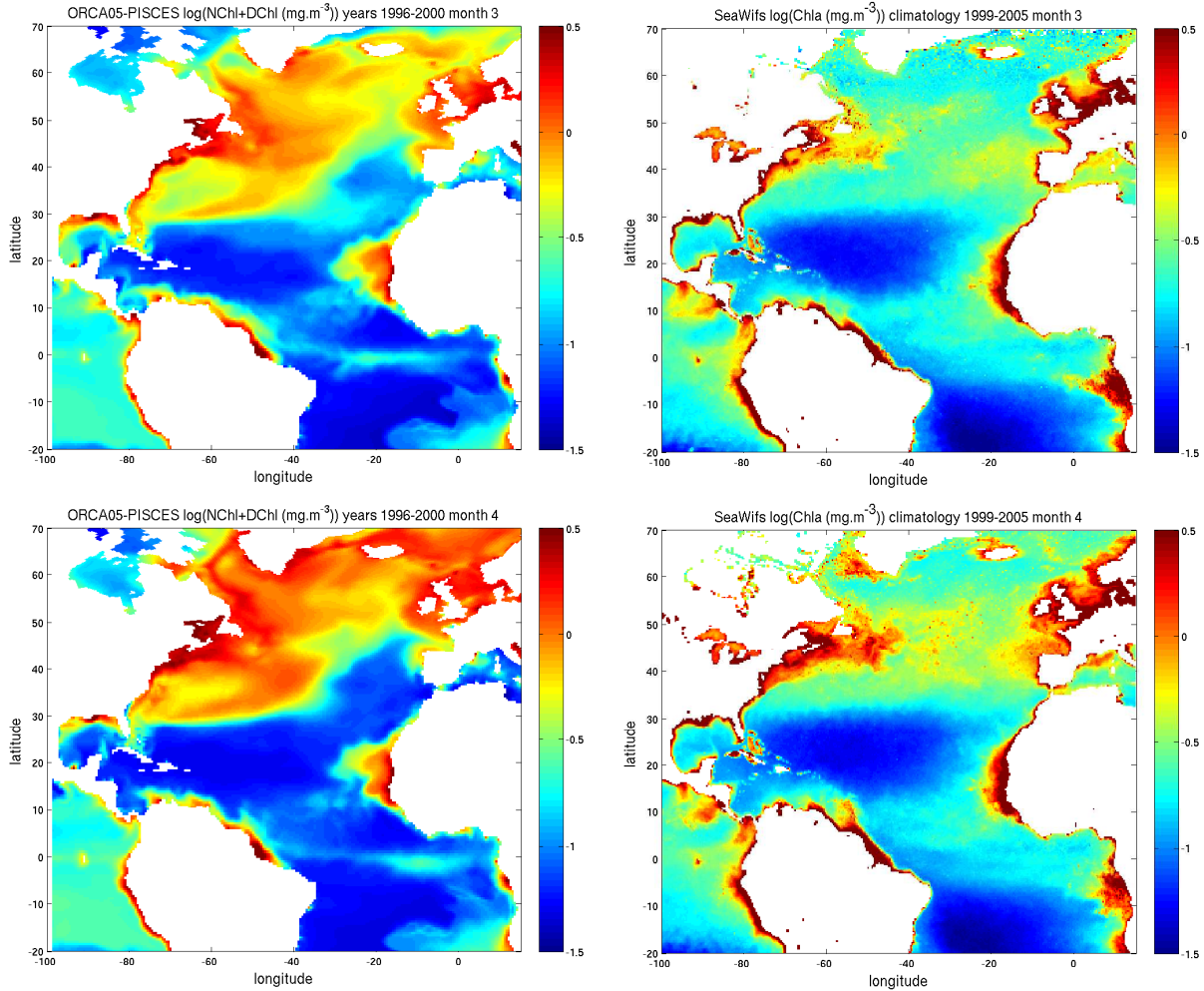


Figure 1: Mean surface chlorophyll in ORCA05-PISCES years 1996-2000 compared with Seawifs climatology in march and april

Chlorophyll concentration is a relevant measurement of the biological activity in the ocean (Figs. 1 and 2, right panels). At mid-latitudes in the North Atlantic, the spring bloom appears north of a latitude around $30^{\circ}N$ - $40^{\circ}N$. In march, april and may, data show a clear meridional gradient of chlorophyll located at a latitude that migrates gradually to the north. In august, high chlorophyll is limited to the subpolar gyre (north of $50^{\circ}N$). The model fails to reproduce this feature. We can see that, although the model has stronger values north of $50^{\circ}N$, there is a permanent deficit in chlorophyll at mid-latitudes in the eastern part of the basin ($30^{\circ}W$ - $10^{\circ}W$ and $30^{\circ}N$ - $50^{\circ}N$). In fact, the North-East Atlantic is completely oligotrophic in the model in the region. between the subtropical and the subpolar gyres. On the other hand, there is too much chlorophyll in the western part of the basin.

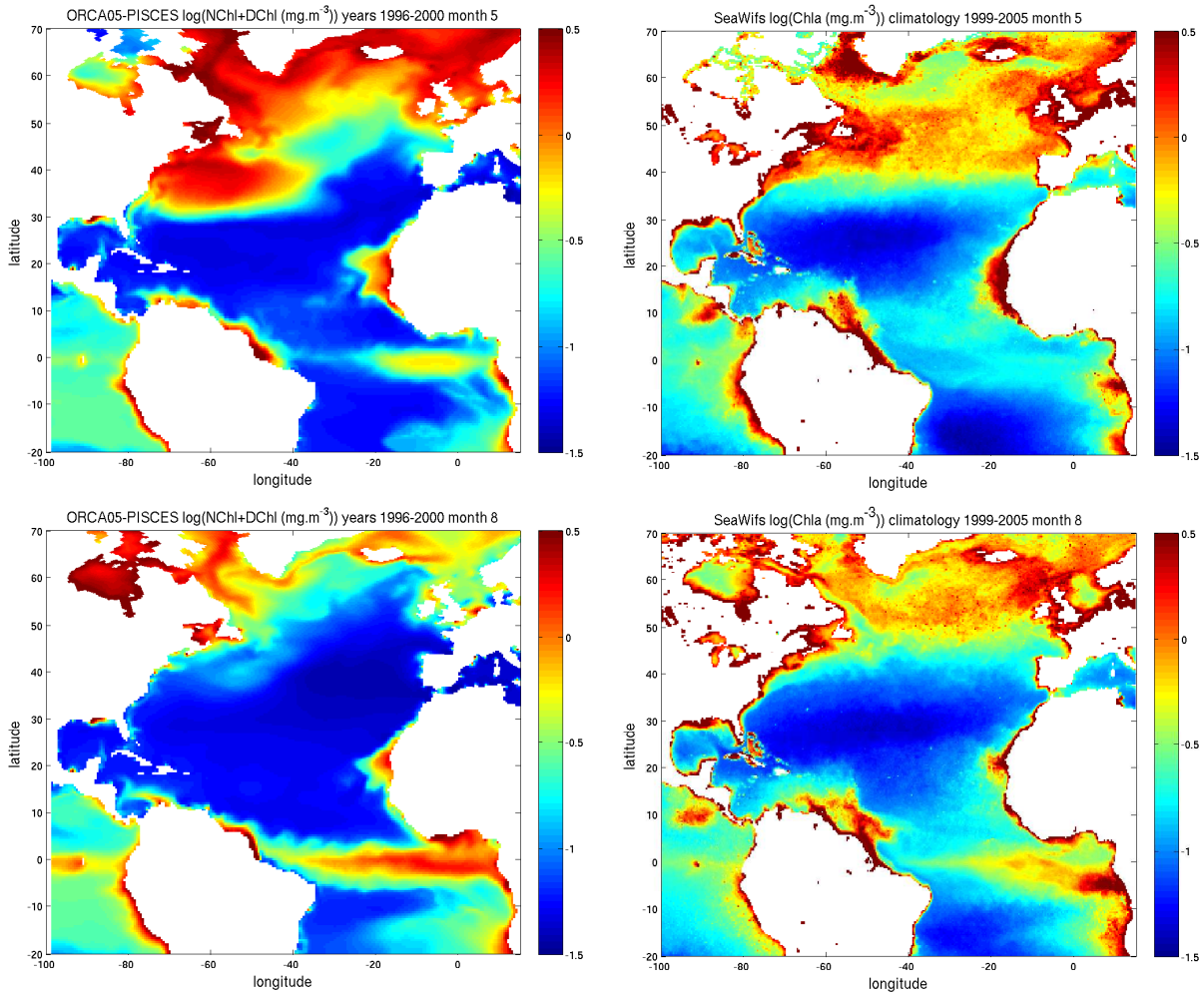


Figure 2: Mean surface chlorophyll in ORCA05-PISCES years 1996-2000 compared with Seawifs climatology in may and august

The model overestimates considerably the chlorophyll gradient from the west (too productive) to the east (too oligotrophic), so we have plotted in Fig.3 zonal sections at 30°N and 40°N to quantify the model-data differences (with a linear scale for chlorophyll). At 30°N observations (blue curves) show a uniform concentration, larger in march and decaying after spring. In contrast the model (red curve) has a larger concentration in the west with an unrealistic east-west gradient. Chlorophyll is always too low in the east. A striking feature is the absence of the concentration maximum (due to the upwelling along the portugal coast) in the model, which contrasts with the good agreement between the model and SEAWIFS in the other great upwelling systems. At 30°N (black curve) observations show large concentrations in the west, a sharp gradient west of 60°W , and then a zonally uniform concentration up to the coastal maximum in the east. The model (magenta curve) creates a spurious gradient in the middle of the basin, with concentrations that fall way below the observed east of 30°W , and again the coastal signal is missing.

The aim of the following figures is to investigate these discrepancies: are they due to the dynamics? If so, is it a problem with the mixed layer depth, the circulation, or with the nutrient supply?

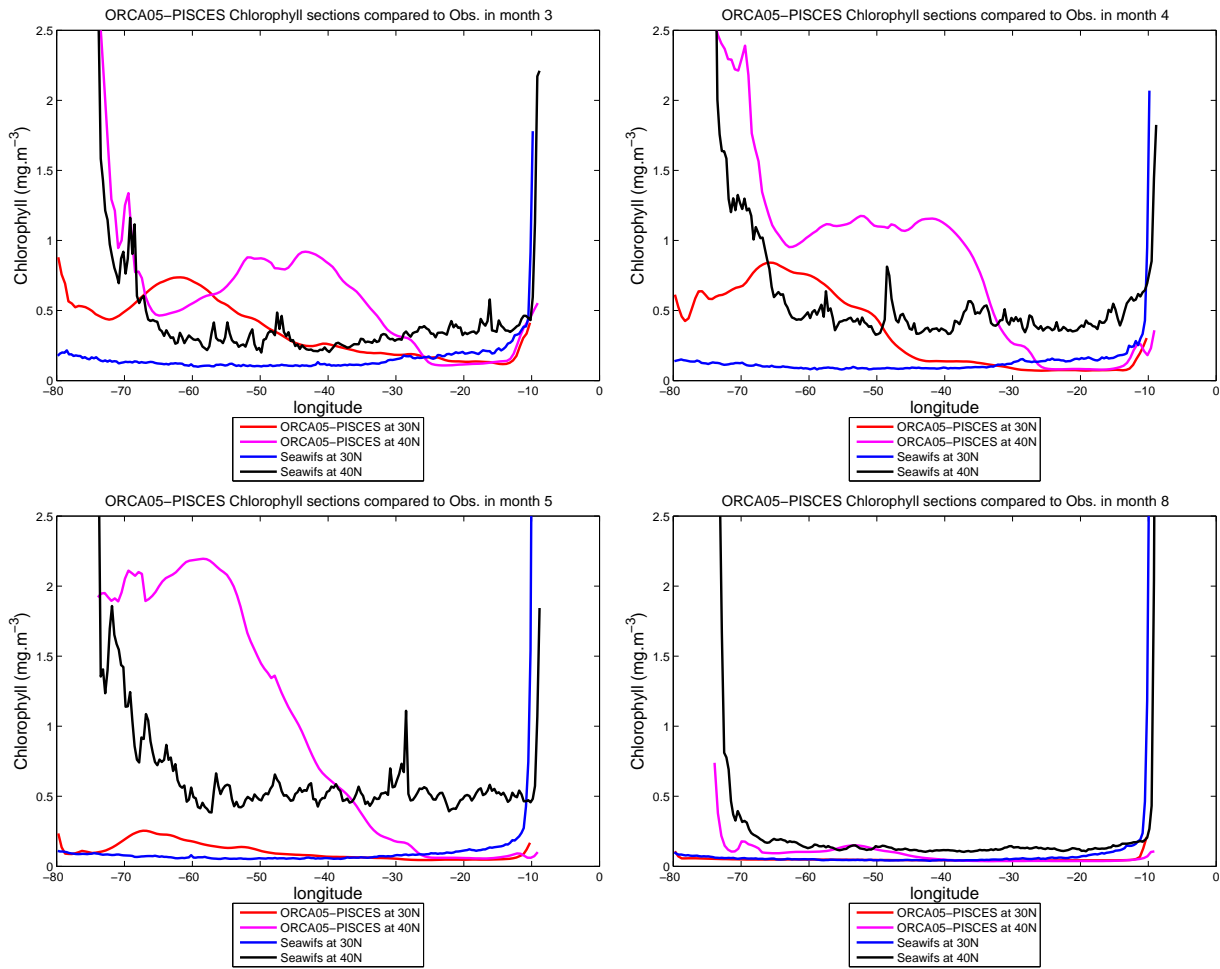


Figure 3: Mean surface chlorophyll in ORCA05-PISCES years 1996-2000 compared with Seawifs climatology in march, april, may and august

3 Comparison of mixed layer depth with Clement de Boyer Montegut climatology

In this section, we compare the mixed layer depth with a climatology based on ARGO floats computed by Clément de Boyer Montégut (de Boyer Montegut et al, ...). Figures 4 and 5 are maps of the mixed layer depth, saturated at 150 meters to emphasize the shallow mixed layers in spring and summer. Figure 6 shows the mixed layer depth along zonal sections at 30 °N and 40 °N in the model and in the climatology, similar to Fig.3.

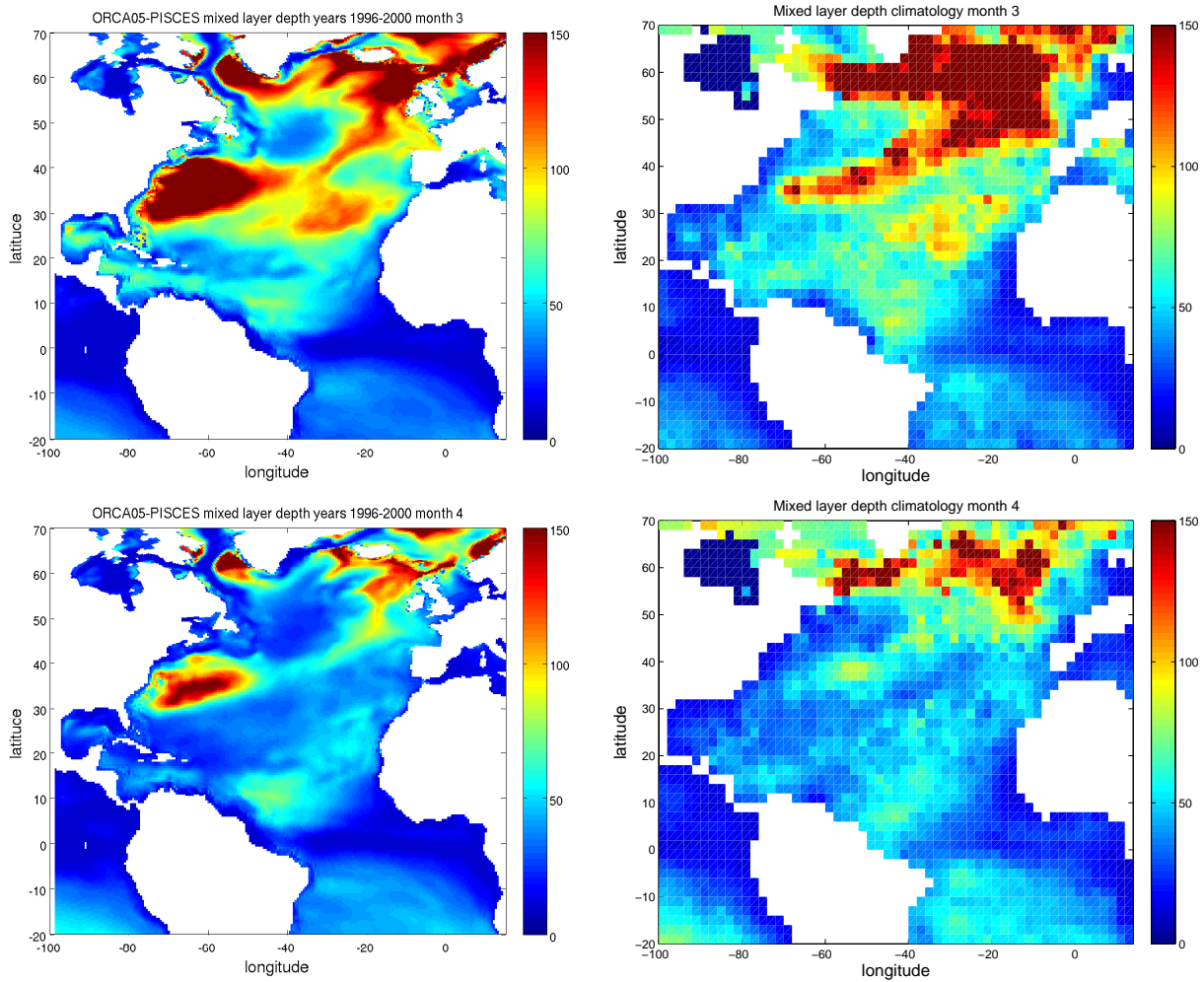


Figure 4: Mixed Layer Depth in ORCA05-PISCES years 1996-2000 compared with climatology in march and april

In march and april, we can see that there is a too deep mixed layer at mid-latitude in the western part of the basin. Values in the model reach 200 meters in march whereas the observation are roughly 50 meters. The reason for this discrepancy could be a misrepresentation of surface fluxes, or the lack of eddies that tend to restratify the mixed layer. In the eastern part of the domain, values in the model have the same order of magnitude than in the observations though some local extrema are not well represented. In may, we have a quite good agreement and the mixed layer is shallower in the model, excepted at 30 °N between 38 °W and 20 °W . In august, the model has almost everywhere a shallower mixed layer than observations, shallower by a factor of roughly 1.5 except in a small area in the west on the 40 °N section.

Generally, the model and observations are in a quite good agreement. Deficiencies of the mixed layer depth do not seem to explain the problem with the primary production.

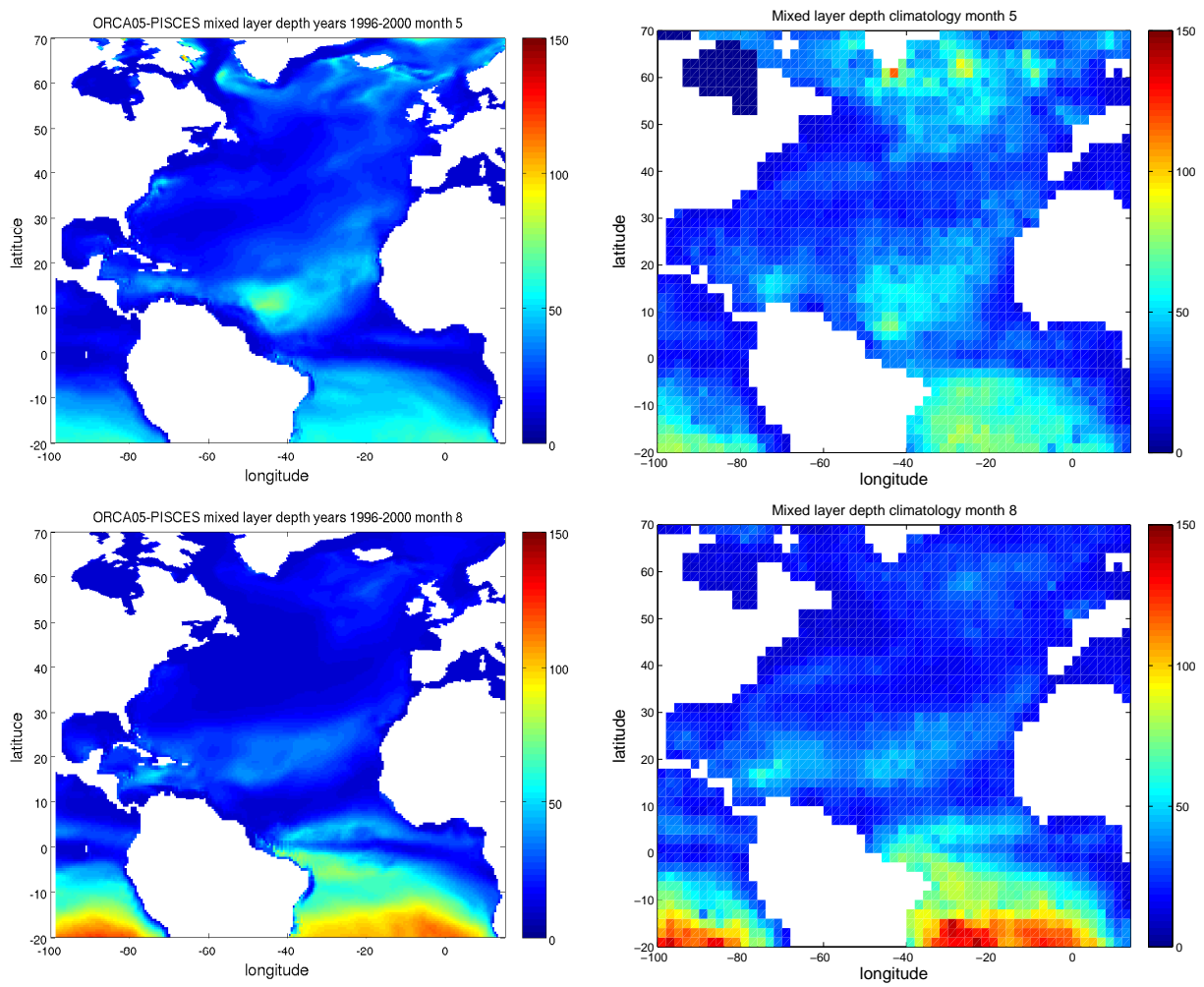


Figure 5: Mixed Layer Depth in ORCA05-PISCES years 1996-2000 compared with climatology in may and august

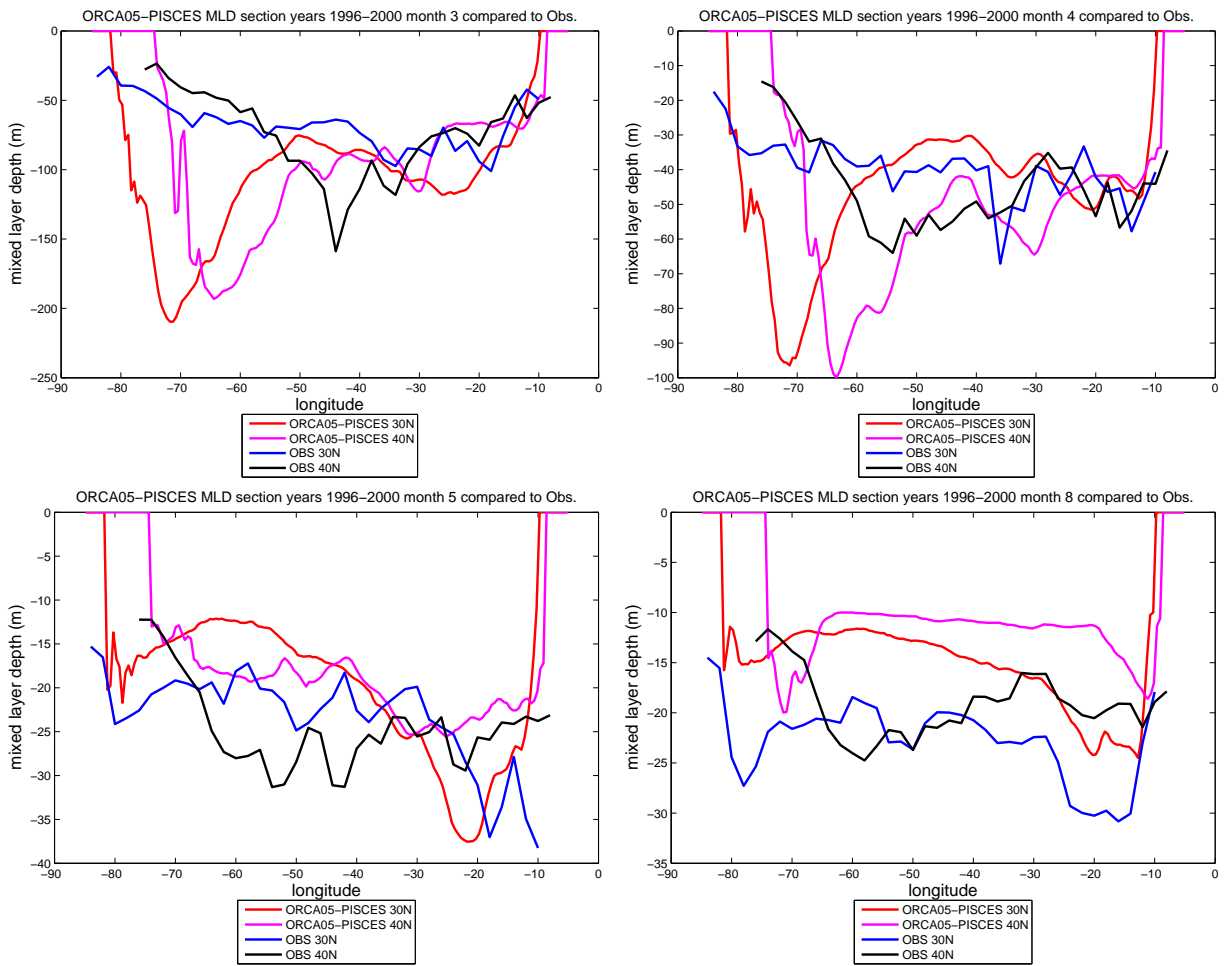


Figure 6: Mixed Layer Depth in ORCA05-PISCES years 1996-2000 compared with climatology in march, april, may and august

4 Comparison of temperature and nitrates sections with World Ocean Atlas

In this section, we compare sections of temperature and nitrates concentration with the World Ocean Atlas at 30°N in march (fig. 7) and august (fig. 8) then at 40°N in march (fig. 9) and august (fig. 10). Figure 11 shows surface nitrates concentration along zonal sections at 30°N and 40°N in the model and in the climatology.

4.1 Section a 30 °N

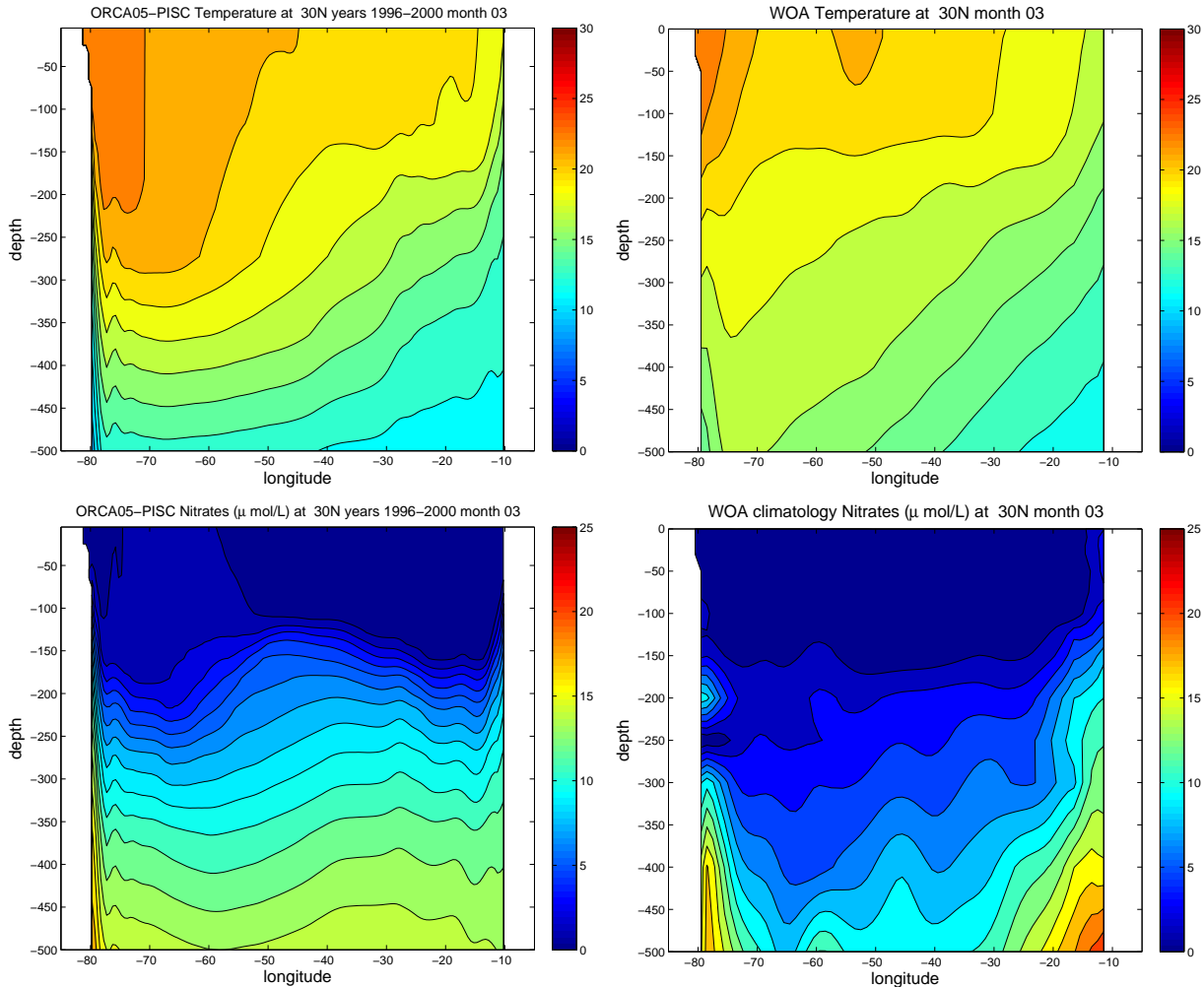


Figure 7: Temperature and Nitrates section at 30 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in march

When we look at the temperature sections at 30 °N , we notice that the extent of hot waters ($T \geq 20$ °C) overshoots by a large amount the location found in the observations and that the thermocline is higher up in the model. The isotherms are too flat which suggest a weak geostrophic flow in the subtropical gyre. The model-data differences in the nitrates sections mirror the differences in temperature: the nitracline is also too flat and shallow and the model miss the deepening in the middle of the basin. The combined discrepancies in nitrates and mixed layer depth could explain the overestimation of the primary production at 30 °N between 65°W and 50°W . The mixed layer depth in march, slightly too deep (200 m) can dig into the nitracline (which is too shallow) and thus provide an unrealistic amount of nutrients. Note that

the seasonal thermocline (restratification in august, fig. 8) is rather well represented by the model, although the mixed layer depth is too shallow (Fig. 6). This may have a consequence on the surface distribution of nitrates as we will show later.

At 40 °N , there are a lot of discrepancies between the model and the observations. The nitrates concentration at depth in the middle of the basin is too high and decrease towards the east whereas the observations show a sharp decrease in the west followed by a more progressive increase. Figure 11 shows that the surface concentration in spring is much larger in the model except in the east or in some small areas and became negligible in august.

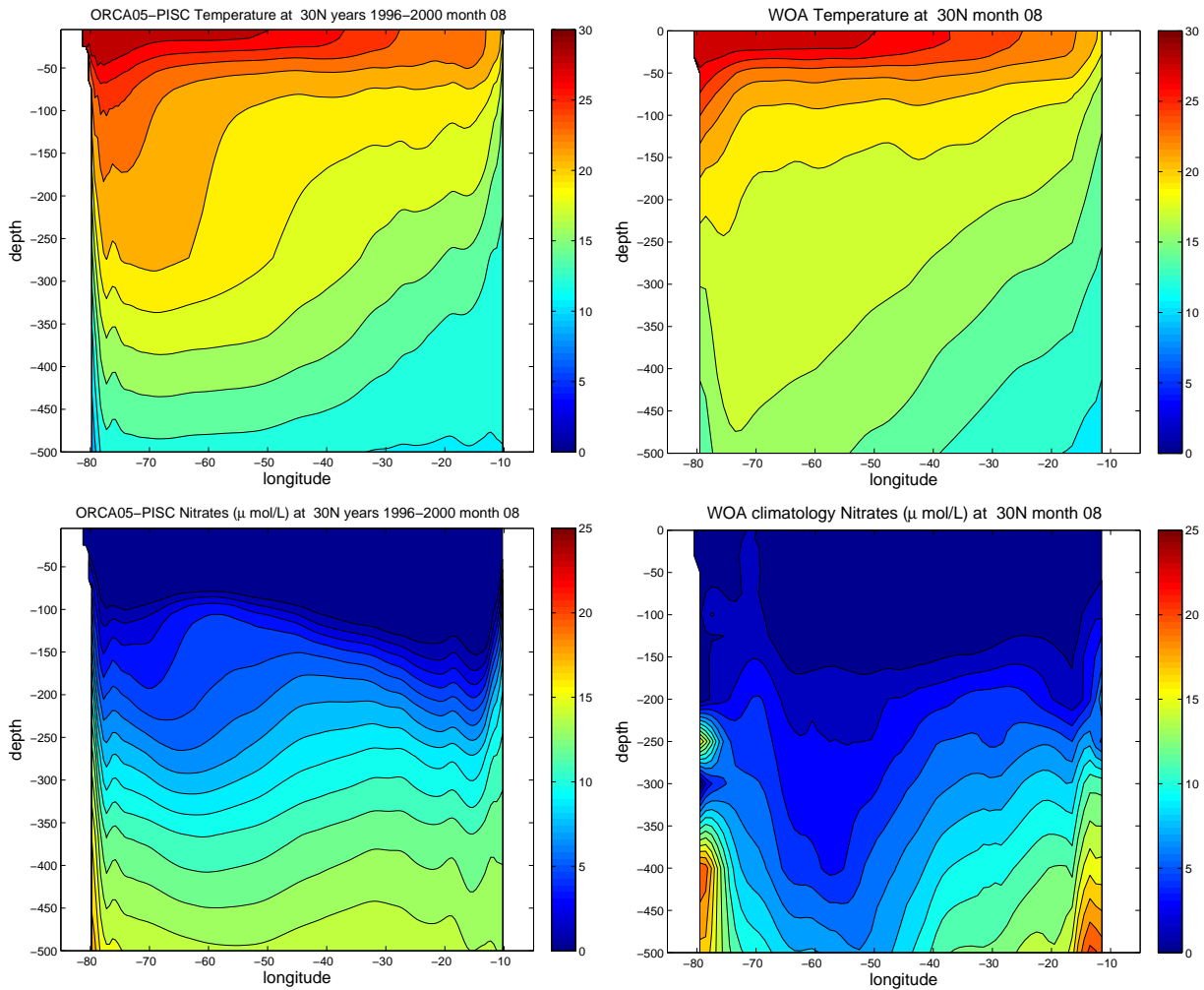


Figure 8: Nitrates section at 30 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in august

4.2 Section a 40 °N

At 40 °N , the model's thermocline structure is also very unrealistic (fig. 9). Again, the thermocline is too flat. The most notable feature is the absence of the sharp gradient near the western boundary, which shows that the strength of the geostrophic shear in the western boundary current is largely underestimated. This is to be expected. The observed transport of the Gulf Stream is about 100 Sv, most of which occurs in the upper layers. The linear Sverdrup circulation is only about 20-30Sv. ORCA05 has no eddies and no inertial recirculation, therefore it represents only the Sverdrup dynamics and cannot reproduce the total baroclinic transport of the Gulf Stream.

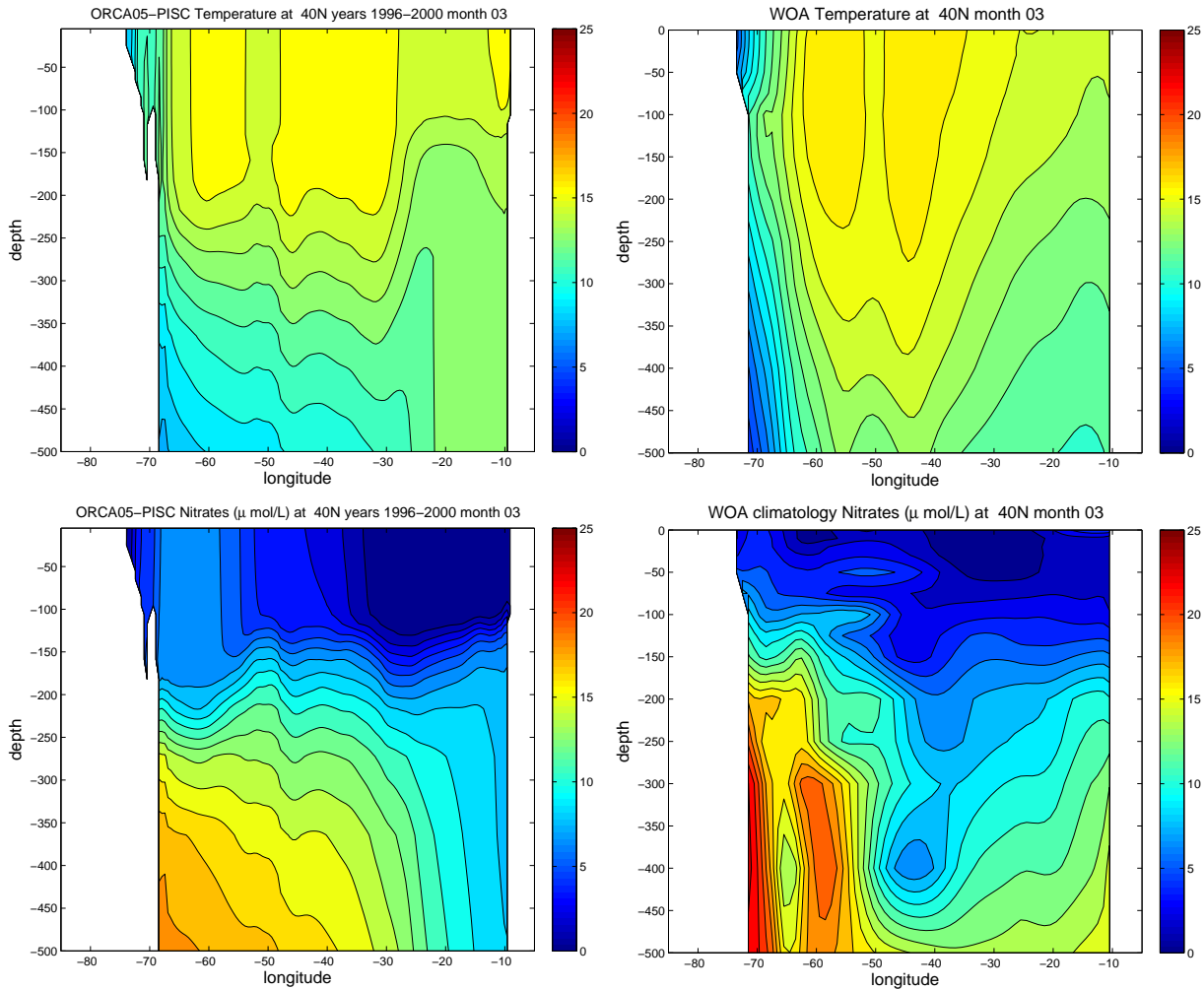


Figure 9: Temperature and Nitrates section at 40 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in march

Moreover, ORCA05 does not represent the cold slope waters advected north of the Gulf Stream from the Labrador shelf. Next to the western boundary, instead of the region of strong gradient between the slope waters and the subtropical waters, the model has no cold water, very weak gradients and a region of vertically homogeneous temperature and nitrates extending down to 200m. Anomalously high levels of nitrates are found near the surface all the way to 40W (fig. 9). This, combined with larger than observed winter mixed layers, is clearly the explanation for the anomalously high primary production west of 40W in the model at 40N. In the eastern part of the section, the model shows a spurious patch of homogeneous waters under 150 meters. This is certainly due to a bad representation of the Mediterranean water overflow, which is too shallow (centered at 400m instead of 800-1000m as observed along the section, see fig 13).

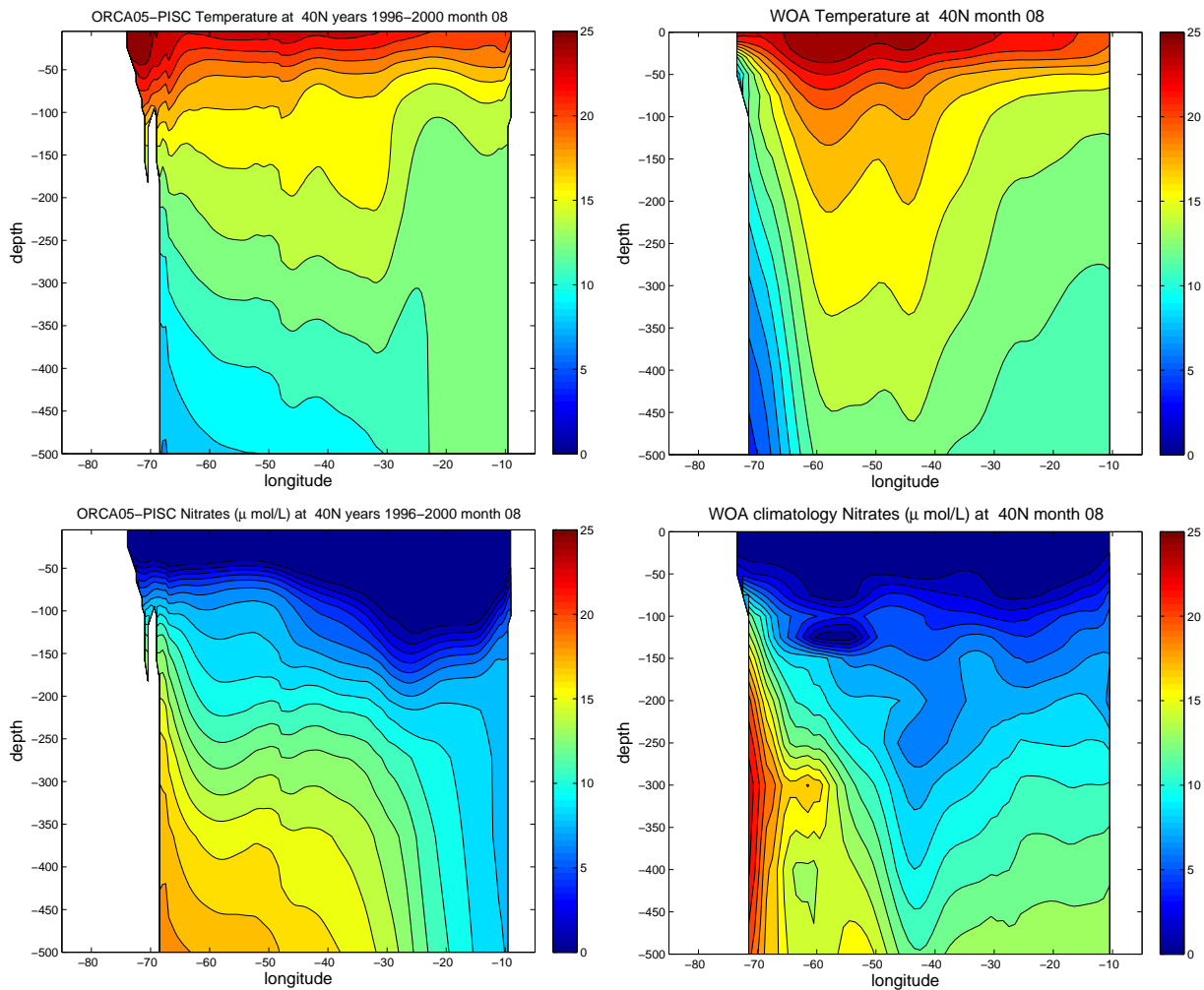


Figure 10: Nitrates section at 40 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in august

Figure 11 shows surface nitrates concentration along zonal sections at 30°N and 40°N in the model and in the climatology. The model considerably underestimates the nitrates in August (why is that? could there be a bug in Levitus? the nitrates are supposed to be exhausted in summer in the mixed layer, are they not??). Could the discrepancy come from the fact that the model mixed layers are too shallow in August?

There are no nitrates at the surface in the model east of 25W, all year round. Why does the model miss completely the source of nitrates in the eastern Atlantic? Is it because of the wrong representation of the thermocline? Or the lack of upwelling along the portuguese coast?

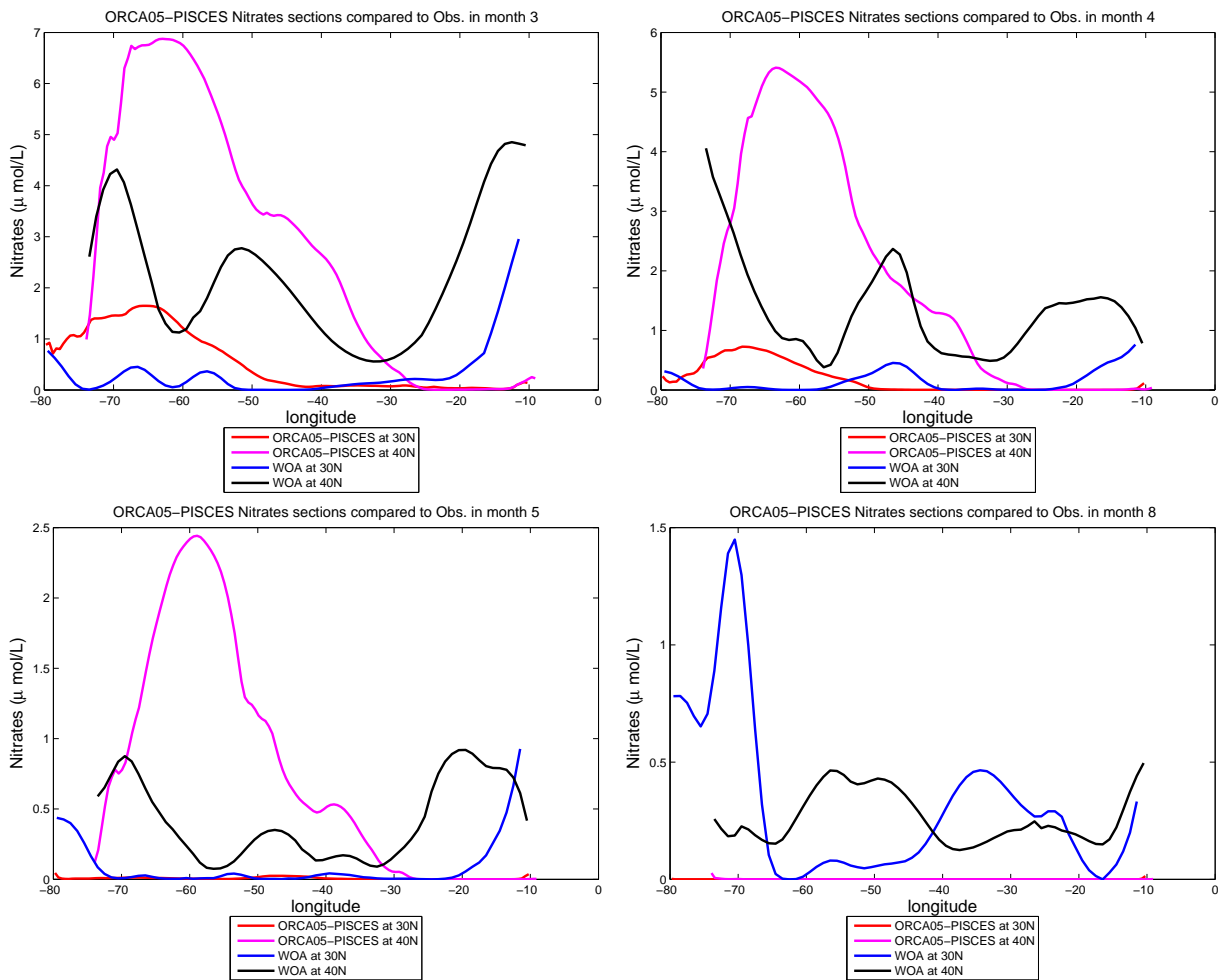


Figure 11: Nitrates sections in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in march, april, may and august

5 Comparison with ORCA05-G70.114

The weakness of the subtropical gyre in ORCA05-PISCINT3 seems to play a part in the bad representation of primary production. In this section, we compare temperature, density and mixed layer depth from ORCA05-PISCINT3 with a DRAKKAR run at the same resolution (namely ORCA05-G70.114) and the World Ocean Atlas. We are not sure about the differences between the two simulations settings, because we don't have all the parameters for the ORCA05-PISCINT3. It seems that a Gent-McWilliams (GM) parameterization is used in ORCA05-PISCINT3 but not in ORCA05-G70.114. The figures show the sections down to 1500 m to better show the bowl-shape structure of the subtropical thermocline. Figures 12 and 13 show a comparison of temperature along section at 30°N and 40°N in march and august, figures 14 and 15 show a comparison of the potential density σ_0 along the same sections and for same months. Finally, figure 16 is a comparison of mixed layer depth between ORCA05-PISCINT3 and ORCA05-G70.114 in march, april, may and august.

Firstly, we look at the temperature sections (cf fig. 12 & 13). At 30°N , we notice that the temperature in ORCA05-G70 is overallly colder than in ORCA05-PISCINT3 which is in better agreement with the World Ocean Atlas. We also have a lower and thicher thermocline which is also an improvement. The flatter "bowl" shape in ORCA05-PISCINT3 could be due the use of the GM parameterization. Furthermore, the mediterranean outflow seems better in ORCA05-G70 because the homogeneous waters at the eastern boundary are better positioned on the vertical. At 40°N (Fig. 13), ORCA05-G70.114 shows a litte improvement in the eastern part of the section but there are very irregular isotherms and the thermocline is too high in the water column. The noisier profiles in the DRAKKAR run could be due to the absence of the GM parameterization.

The same features appear in the density section (cf fig. 14 & 15). At 30°N , density in ORCA05-PISCINT3 and ORCA05-G70.114 is very similar but the latter shows a tiny improvement on the slope of isopycnets. At 40°N , we have an improvement at subsurface but the depth and the thickness of the pycnocline are very similar. In conclusion, the DRAKKAR run ORCA05-G70.114 represents only a small improvement over ORCA05-PISCINT3, and it is quite likely that running pisces with the DRAKKAR version of ORCA05 would not improve considerably the biogeochemistry.

5.1 Temperature section at 30 °N

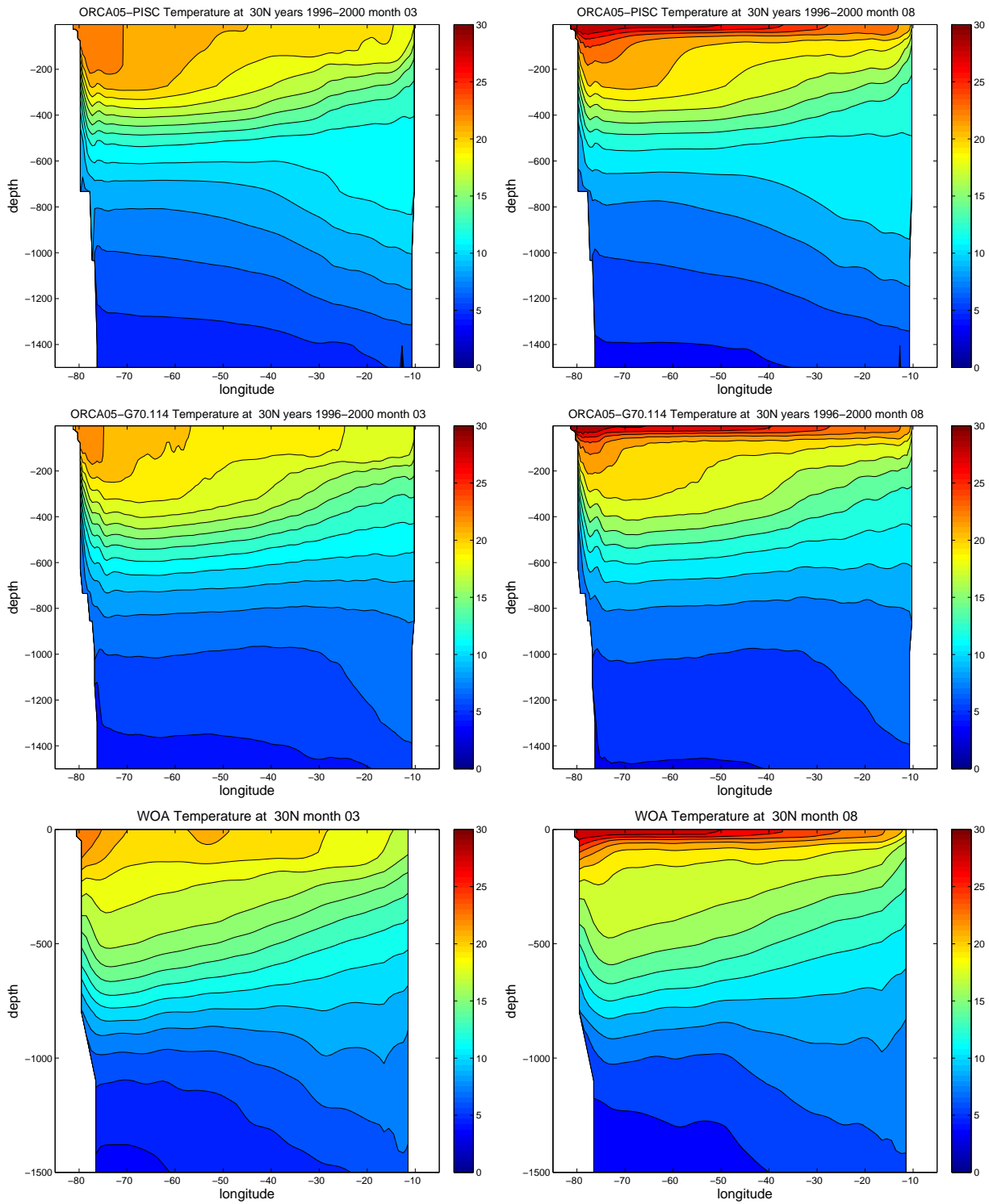


Figure 12: Temperature section at 30 °N in ORCA05-PISCES years 1996-2000 compared with ORCA05-G70.114 and World Ocean Atlas in march and august

5.2 Temperature section at 40 °N

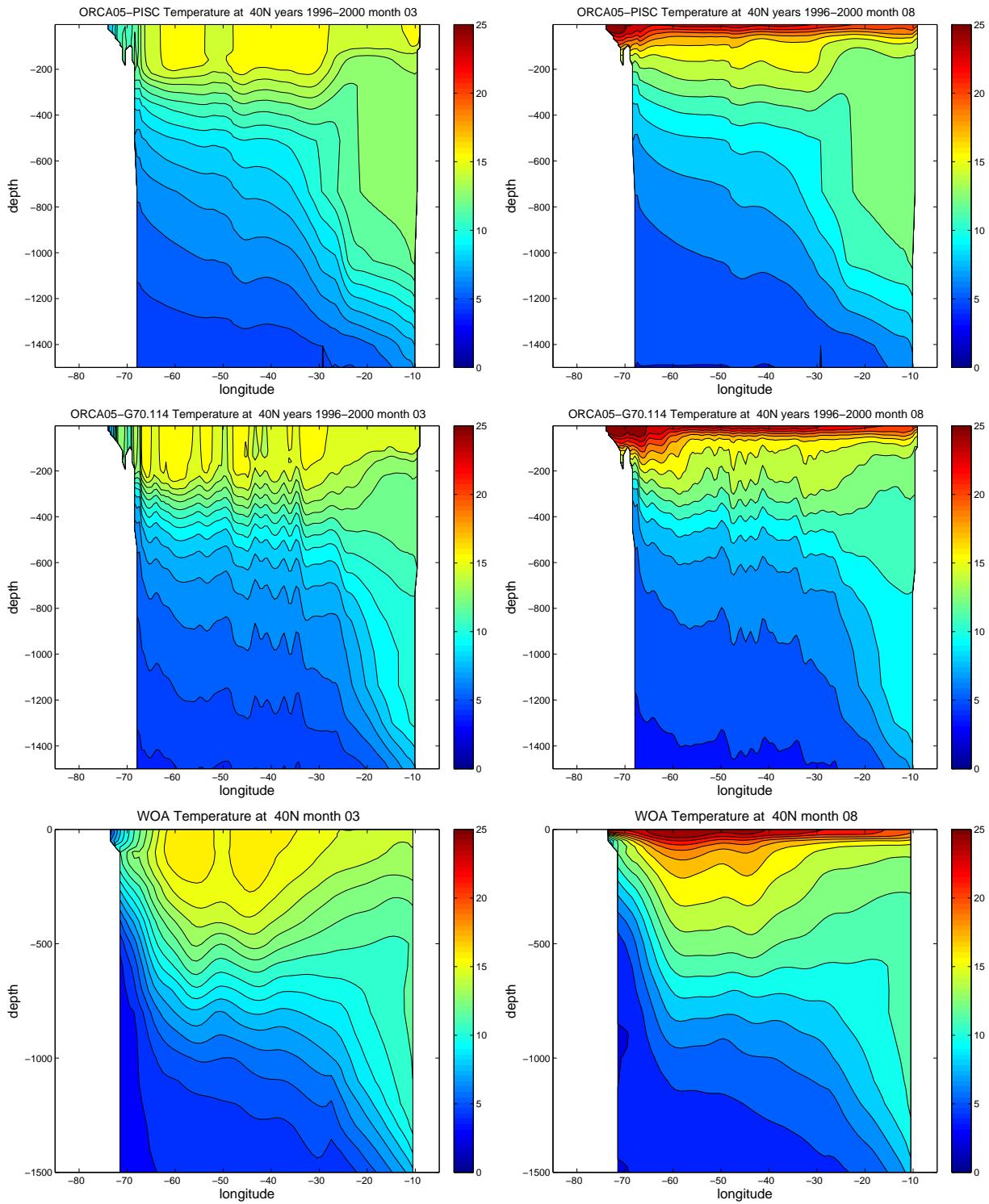


Figure 13: Temperature section at 40 °N in ORCA05-PISCES years 1996-2000 compared with ORCA05-G70.114 and World Ocean Atlas in march and august

5.3 Density section at 30 °N

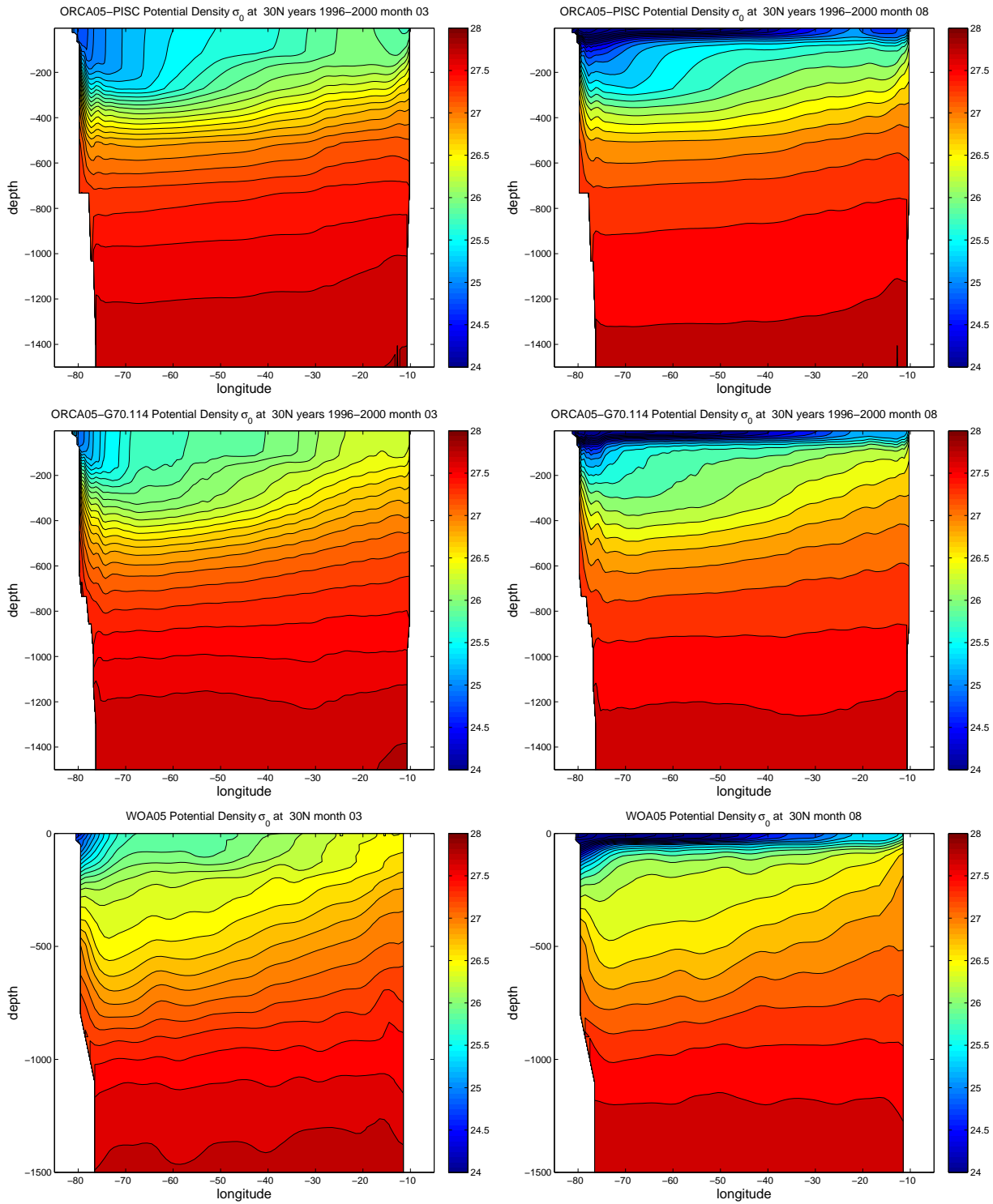


Figure 14: Density section at 30 °N in ORCA05-PISCES years 1996-2000 compared with ORCA05-G70.114 and World Ocean Atlas in march and august

5.4 Density section at 40 °N

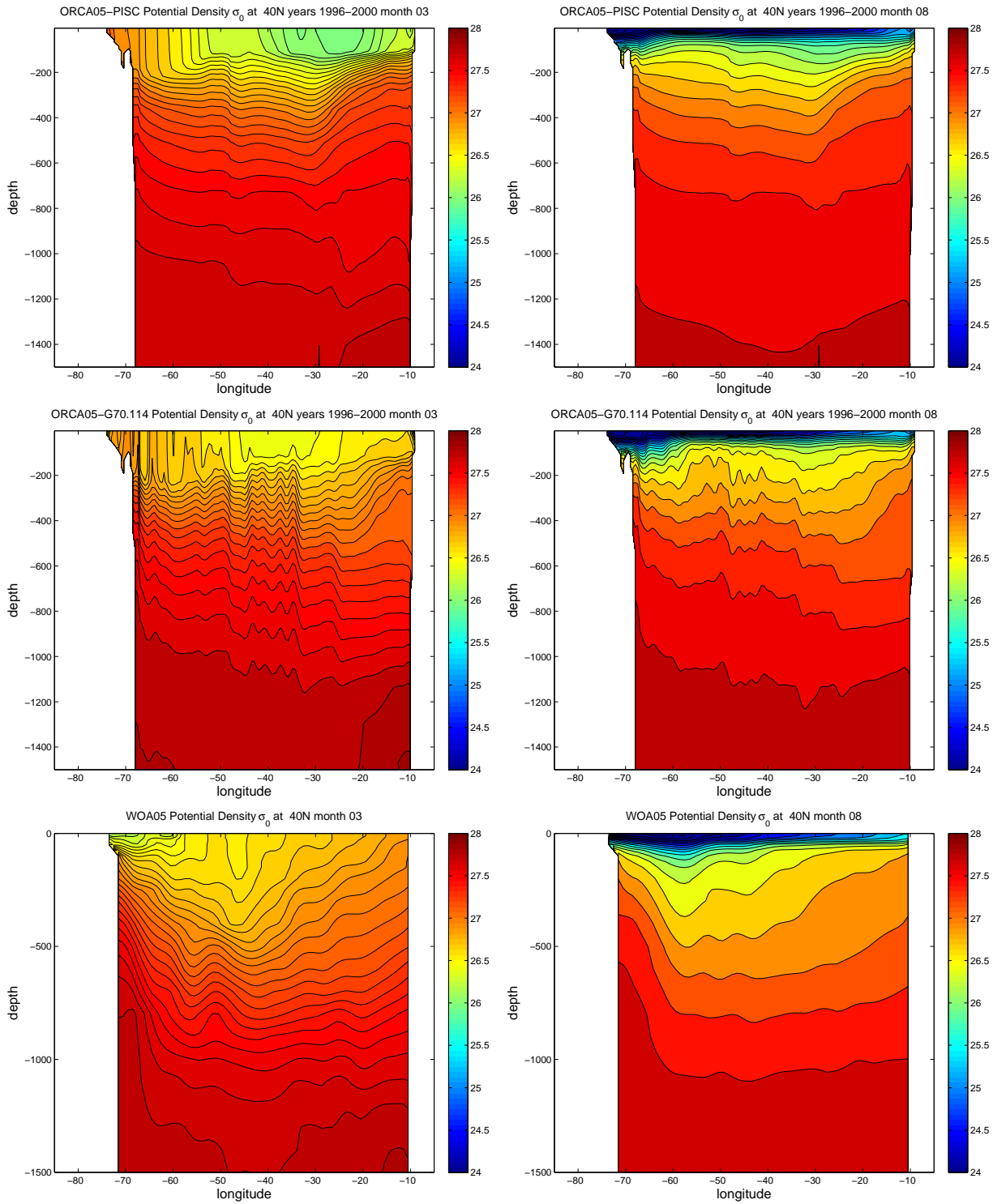


Figure 15: Density section at 40 °N in ORCA05-PISCES years 1996-2000 compared with ORCA05-G70.114 and World Ocean Atlas in march and august

5.5 Mixed layer depth sections at 30 and 40°N

At 30°N , ORCA05-G70.114 has a shallower mixed layer in the western part of the section in spring which is an improvement compared to ORCA05-PISCINT3. In the remaining part of the section, the mixed layer of ORCA05-G70.114 is comparable to ORCA05-PISCINT3 but the former is slightly shallower especially in the eastern part. At 40°N , we notice that the mixed layer of ORCA05-G70.114 is deeper in april and may between 70°W and 50°W , slightly deeper in the east in spring but globally very comparable with ORCA05-PISCINT3.

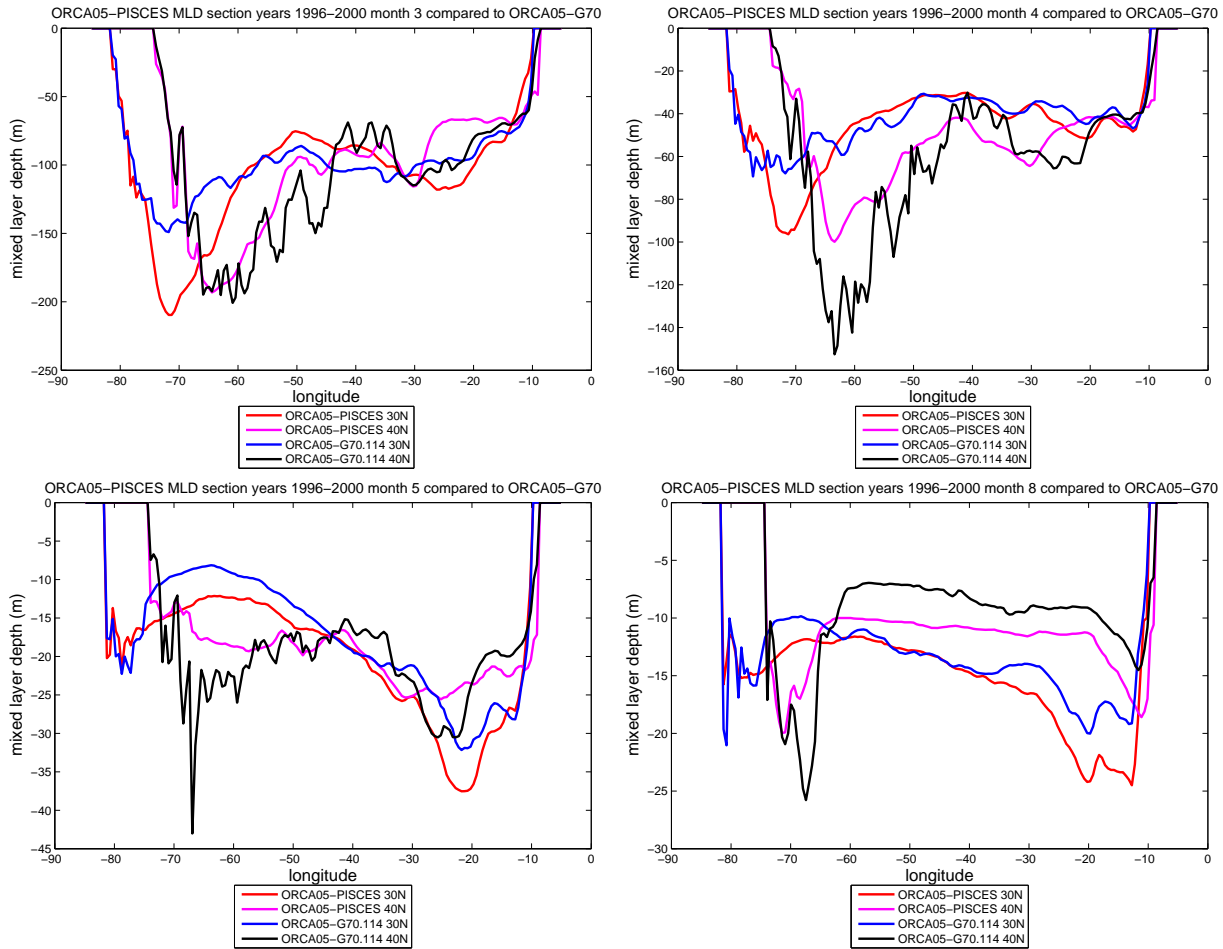


Figure 16: Mixed Layer Depth in ORCA05-PISCES years 1996-2000 compared with ORCA05-G70.114 in march, april, may and august

6 Results from ORCA025-G70 and NATL12-BAMT20

Let us now compare the ORCA05 results with DRAKKAR simulations at higher resolution. We use ORCA025-G70, a 1/4 degree global run (Molines et al, 2006) and NATL12-BAMT20, a north Atlantic 1/12 simulation (Treguier 2008).

The shape of the pycnocline and the representation of Mode waters seem to improve with resolution. NATL12 provides the best solution.

6.1 Temperature section at 30 °N

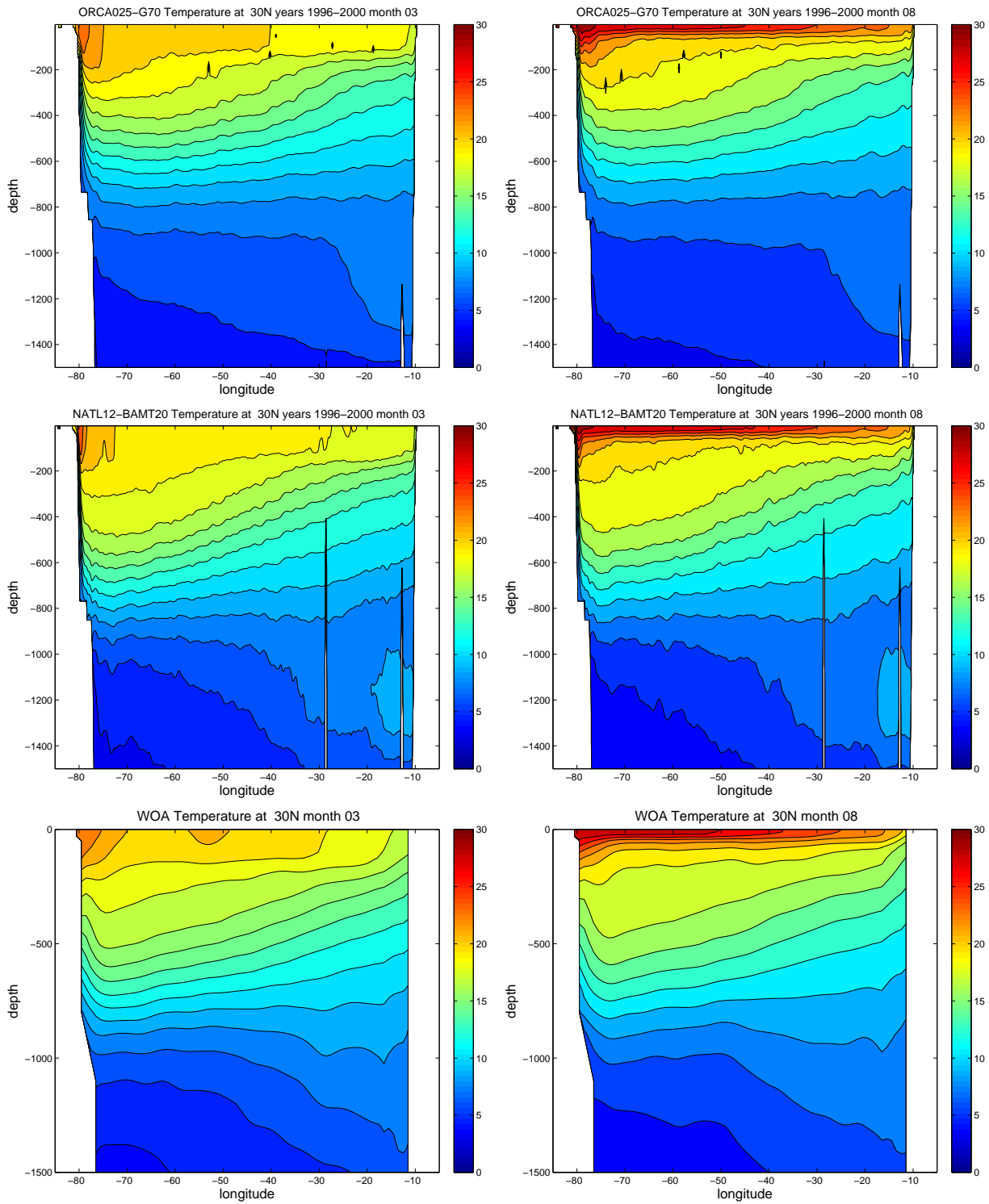


Figure 17: Temperature section at 30 °N in ORCA025-G70, NATL12-BAMT20 years 1996-2000 compared with World Ocean Atlas in march and august

6.2 Temperature section at 40 °N

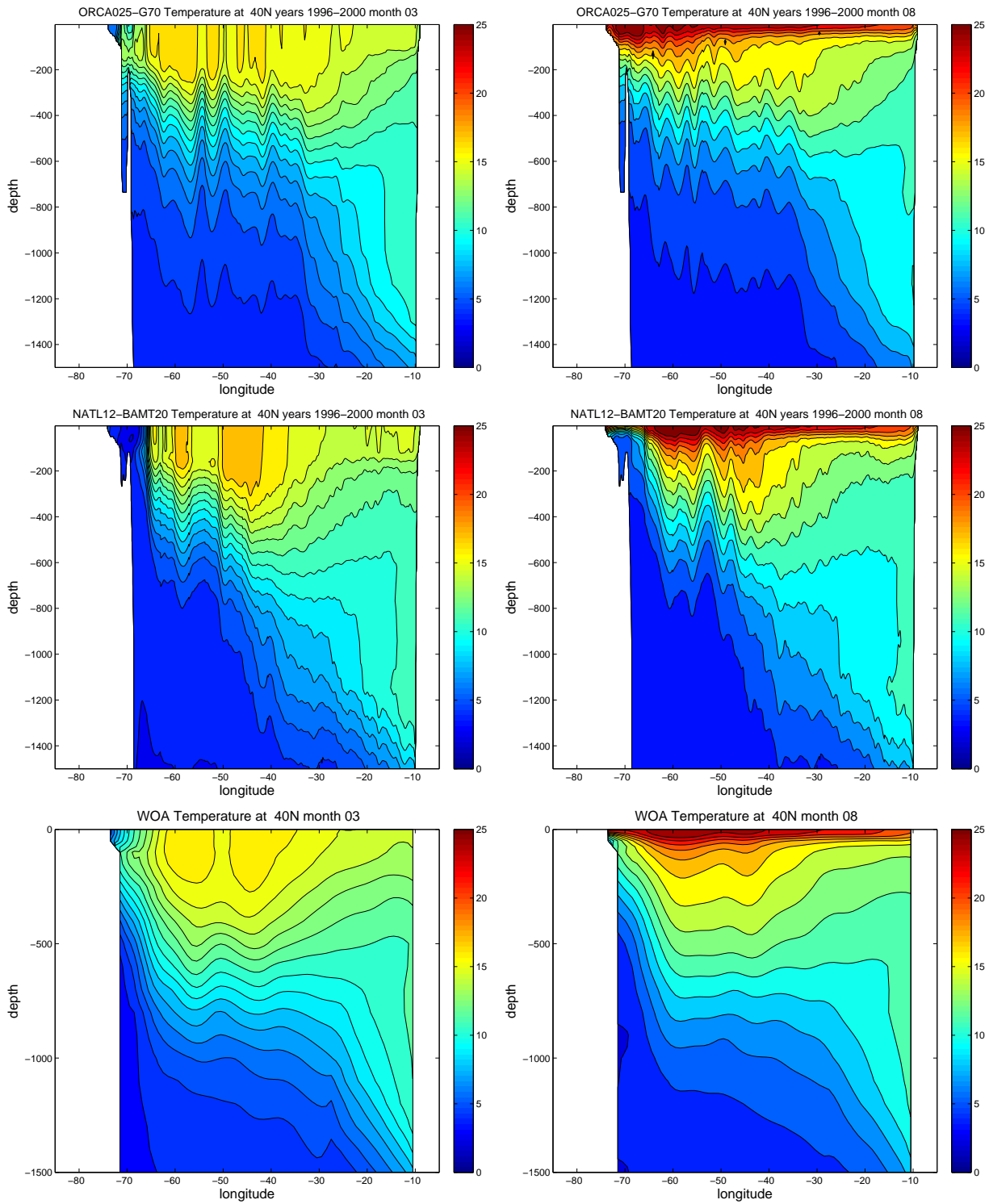


Figure 18: Temperature section at 40 °N in ORCA025-G70, NATL12-BAMT20 years 1996-2000 compared with World Ocean Atlas in march and august

6.3 Density section at 30 °N

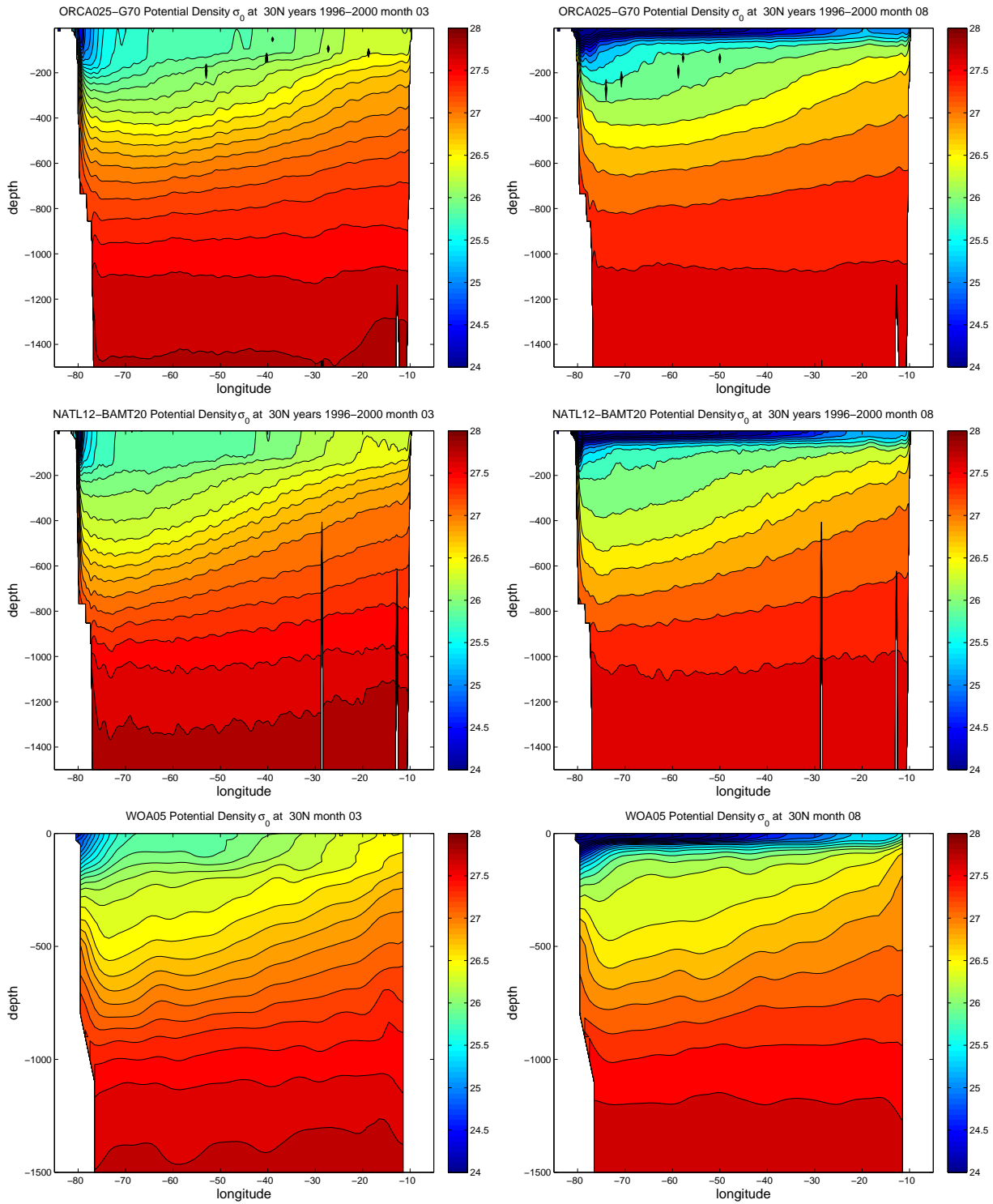


Figure 19: Density section at 30 °N in ORCA025-G70, NATL12-BAMT20 years 1996-2000 compared with World Ocean Atlas in march and august

6.4 Density section at 40 °N

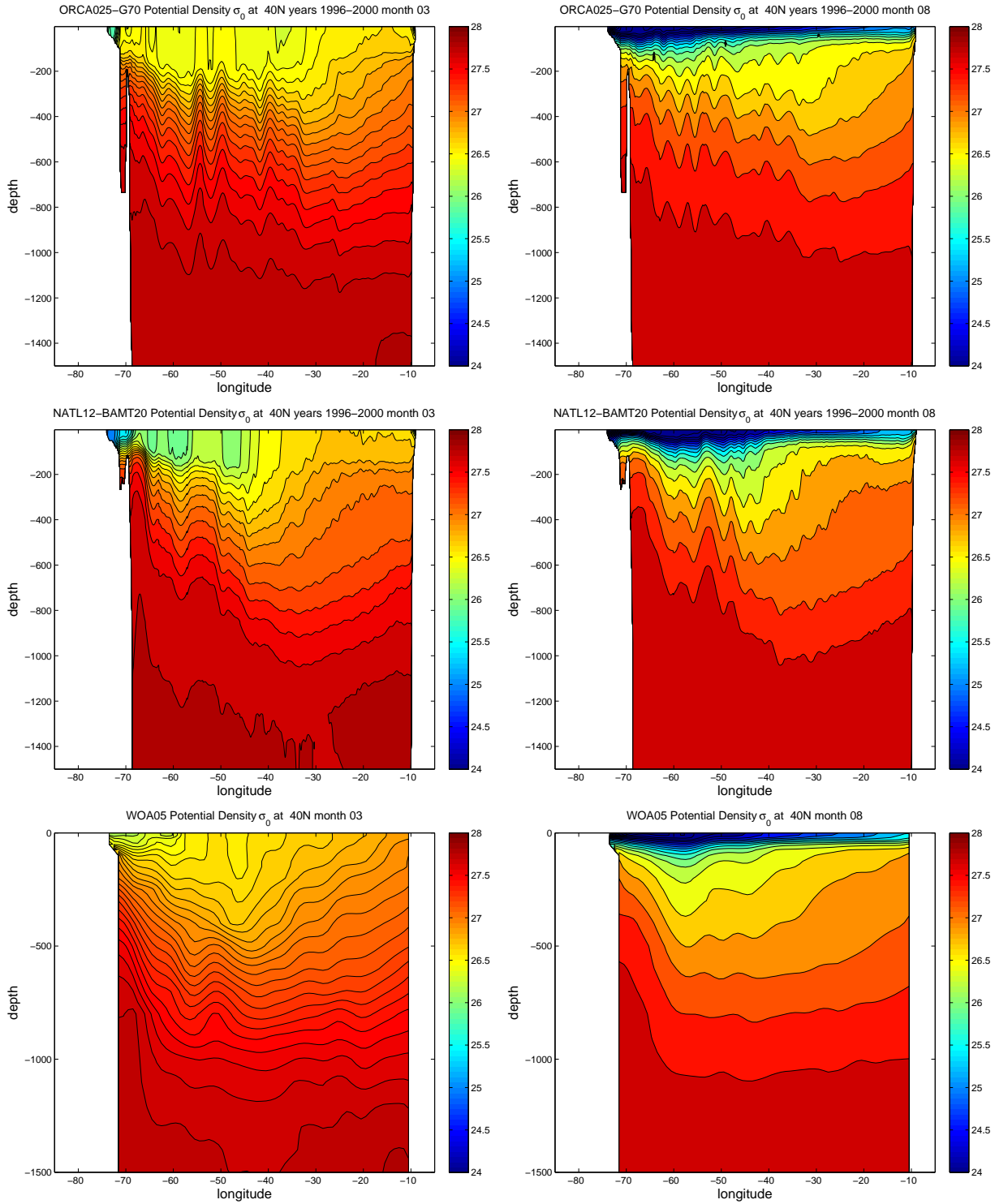


Figure 20: Density section at 40 °N in ORCA025-G70, NATL12-BAMT20 years 1996-2000 compared with World Ocean Atlas in march and august

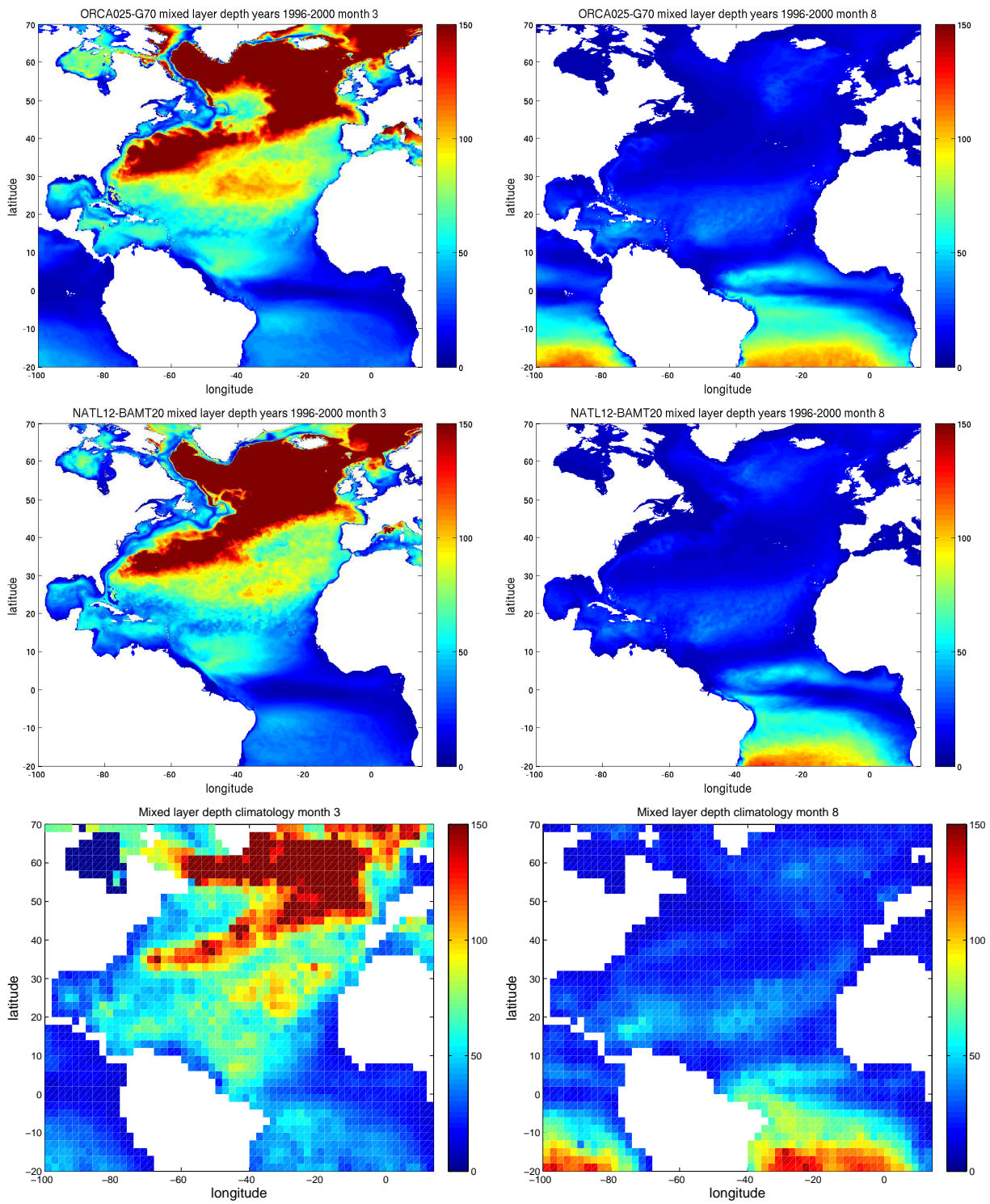


Figure 21: Mixed layer depth in ORCA025-G70, NATL12-BAMT20 years 1996-2000 compared with Climatology in march and august

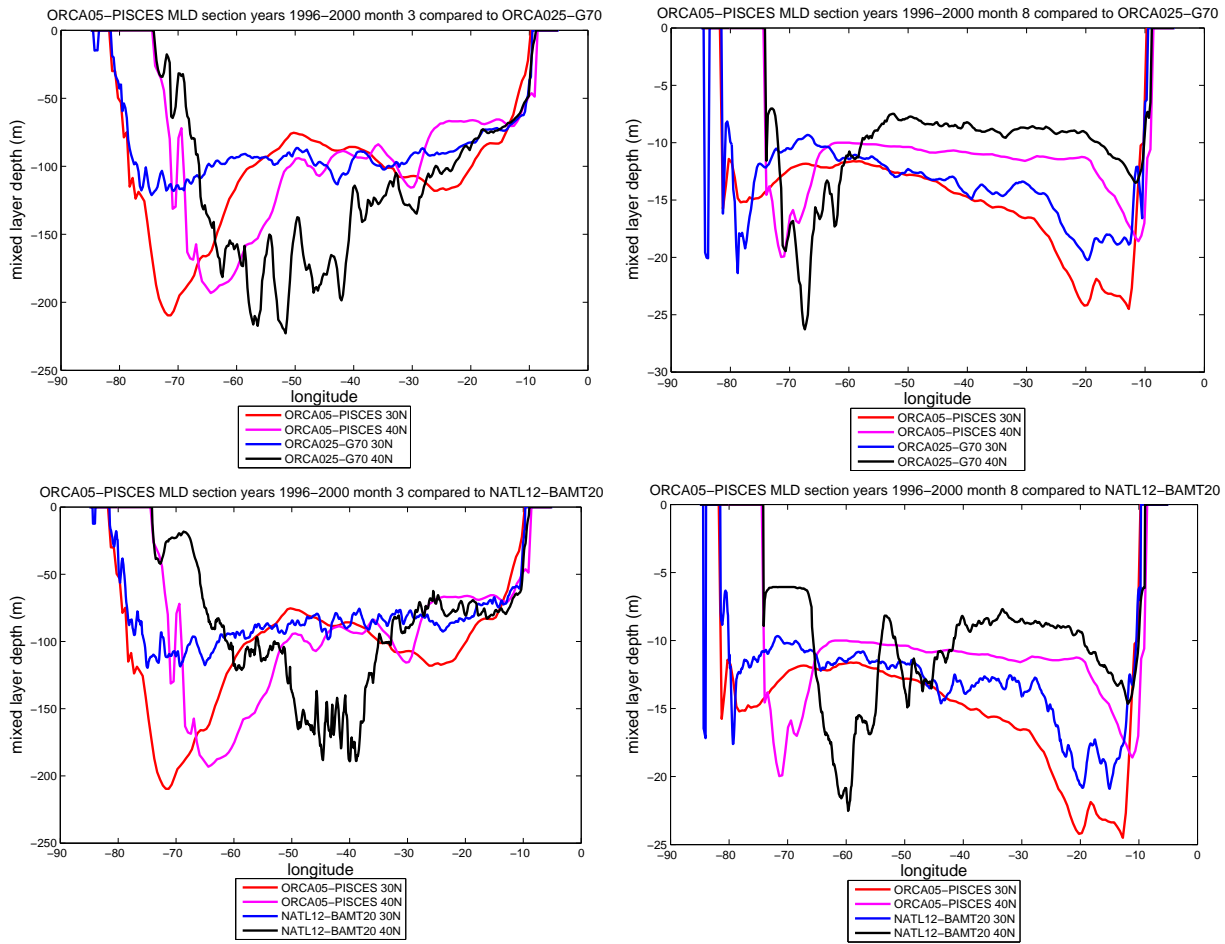


Figure 22: Mixed Layer Depth in ORCA05-PISCES years 1996-2000 compared with ORCA025-G70 and NATL12-BAMT20 in march and august

Appendices

Annex 1 : results from Mercator ORCA025 offline experiment

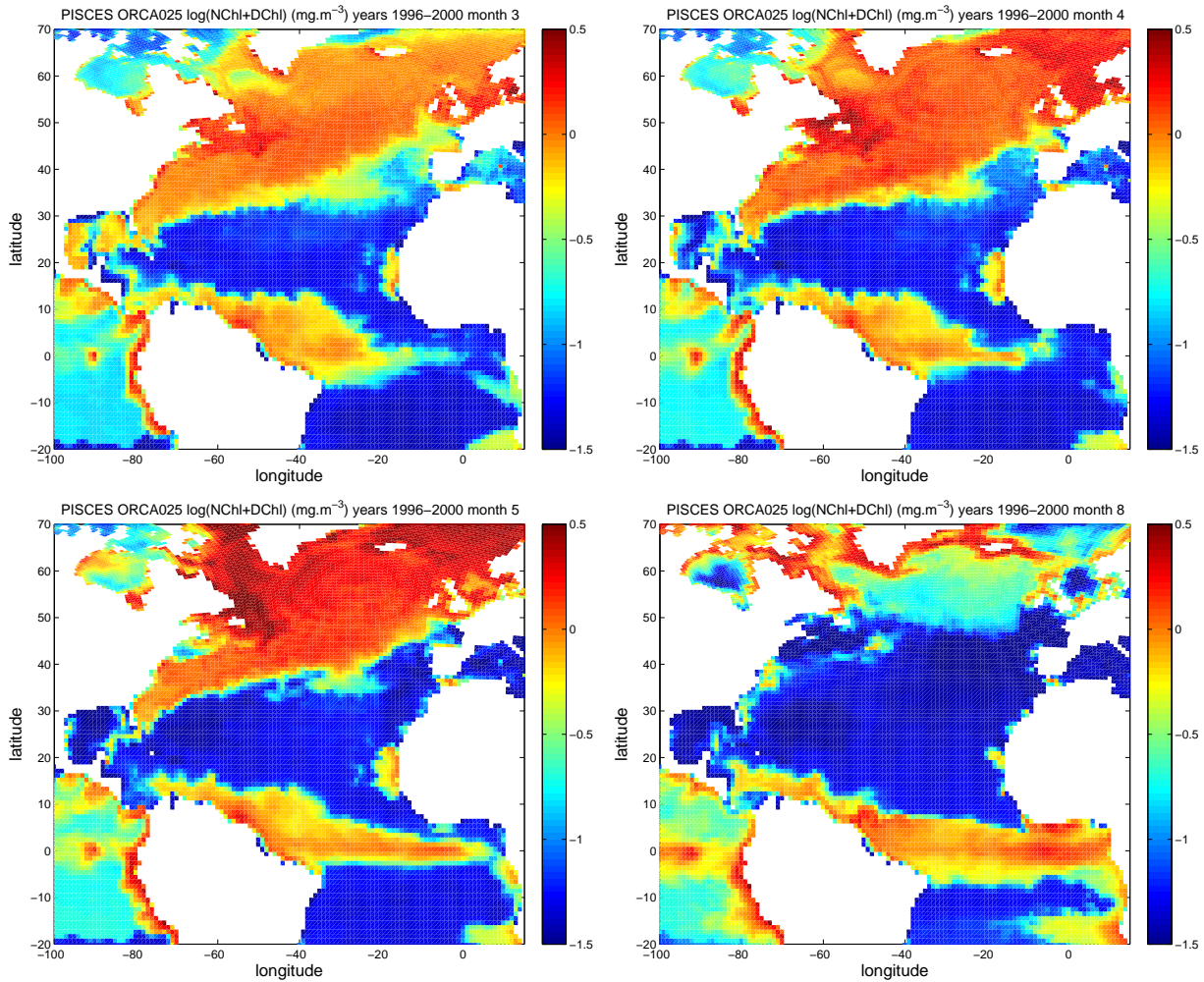


Figure 23: Mean surface chlorophyll in PISCES-ORCA025-BIO1-DMP3 year 2002 march, april, may and august

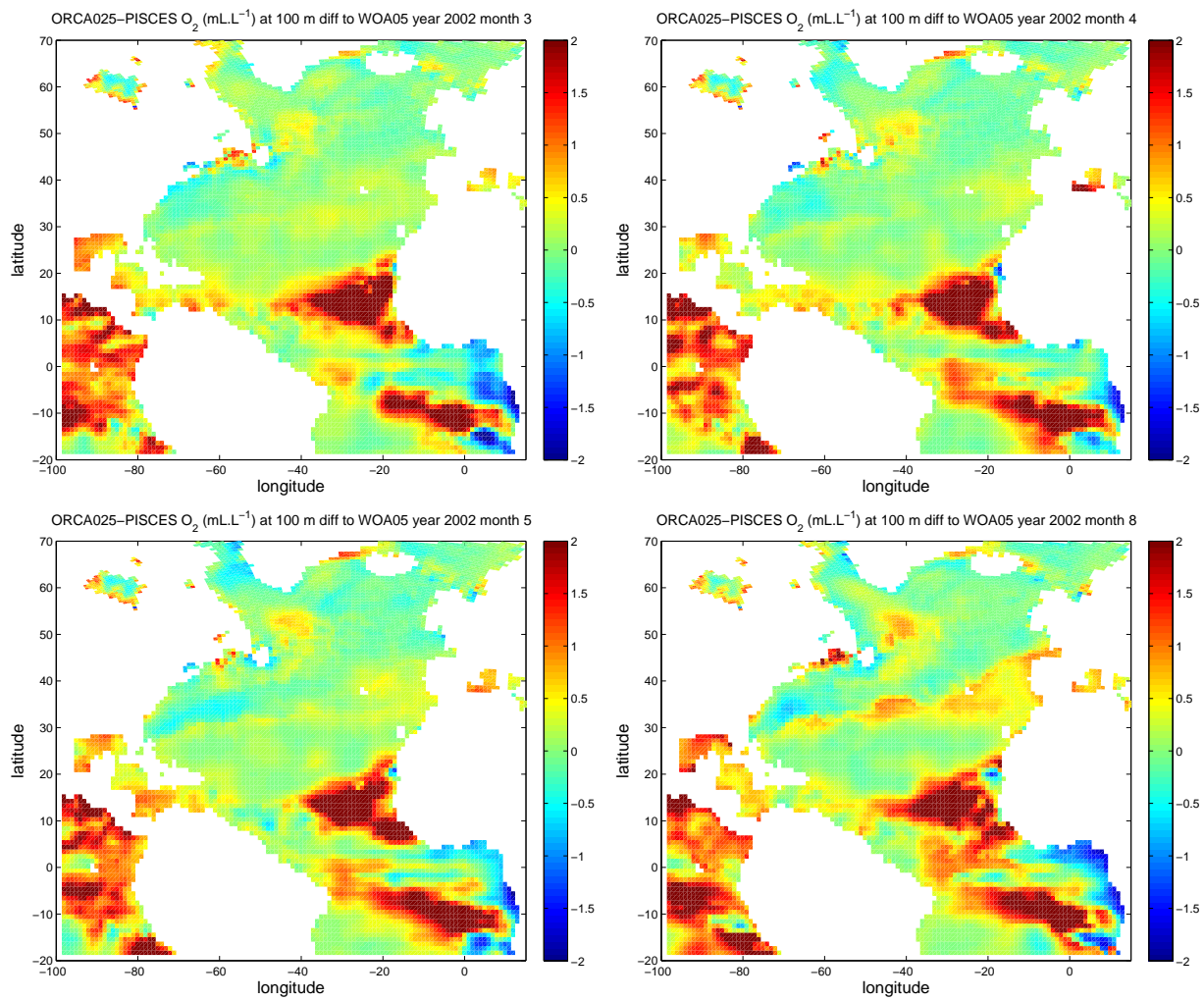


Figure 24: Mean oxygen at 100m in ORCA025-PISCES year 2002 diff to WOA05 in march, april, may and august

Annex 2 : Comparaison of Temperature and Salinity with World Ocean Atlas

We compute the difference between the model and the World Ocean Atlas 2005 climatology. Looking at the temperature maps, we see that the model is too hot in most of the coastal regions (along Africa, North and South America) and in the tropics. It is also too cold in the Labrador Sea and below. The salinity maps shows has a minimum of salinity in the middle of the Atlantic at 45 °N and an excess of salinity in the vicinity of major rivers.

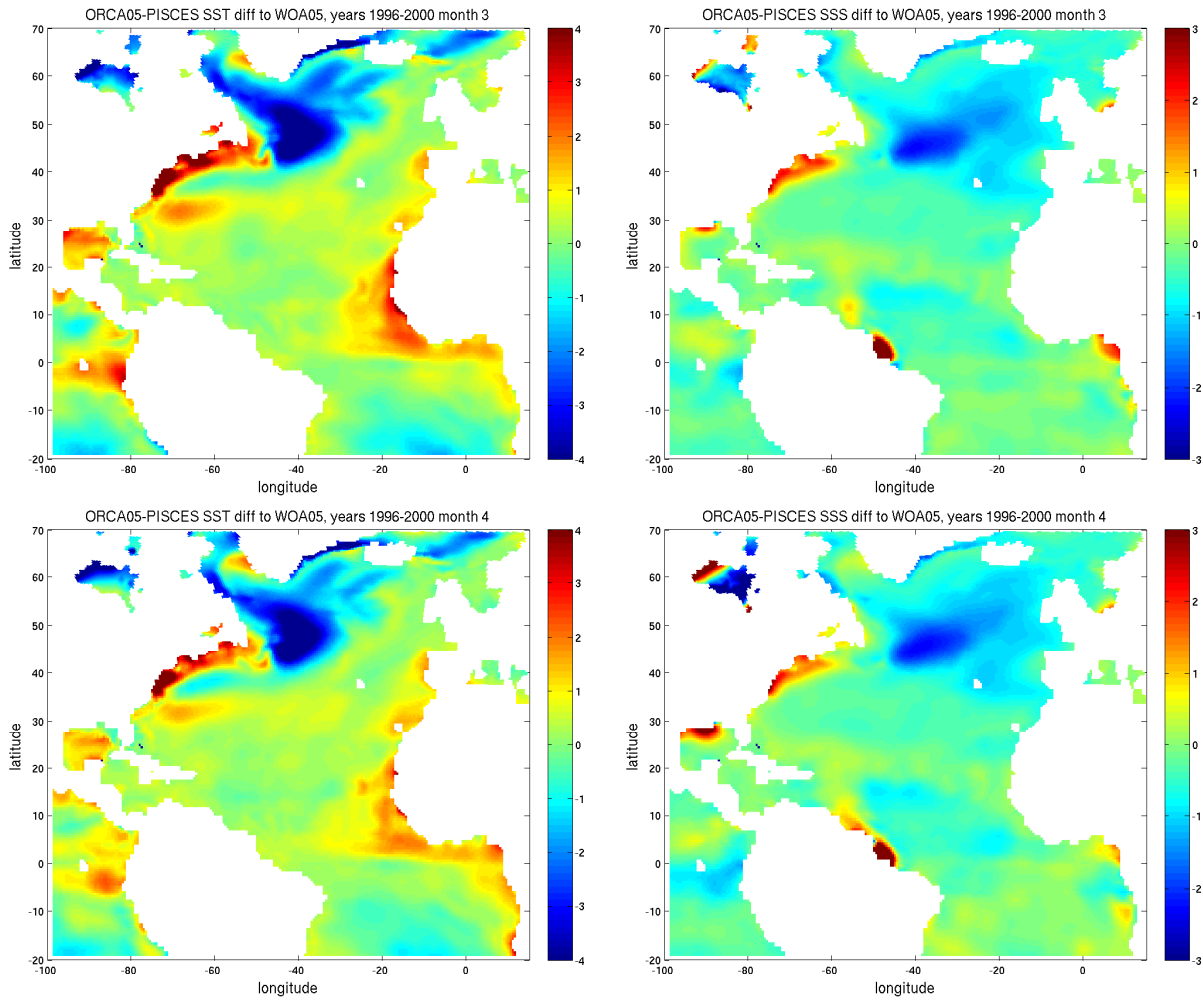


Figure 25: SST & SSS : ORCA05-PISCES diff to WOA05, years 1996-2000 in march and april

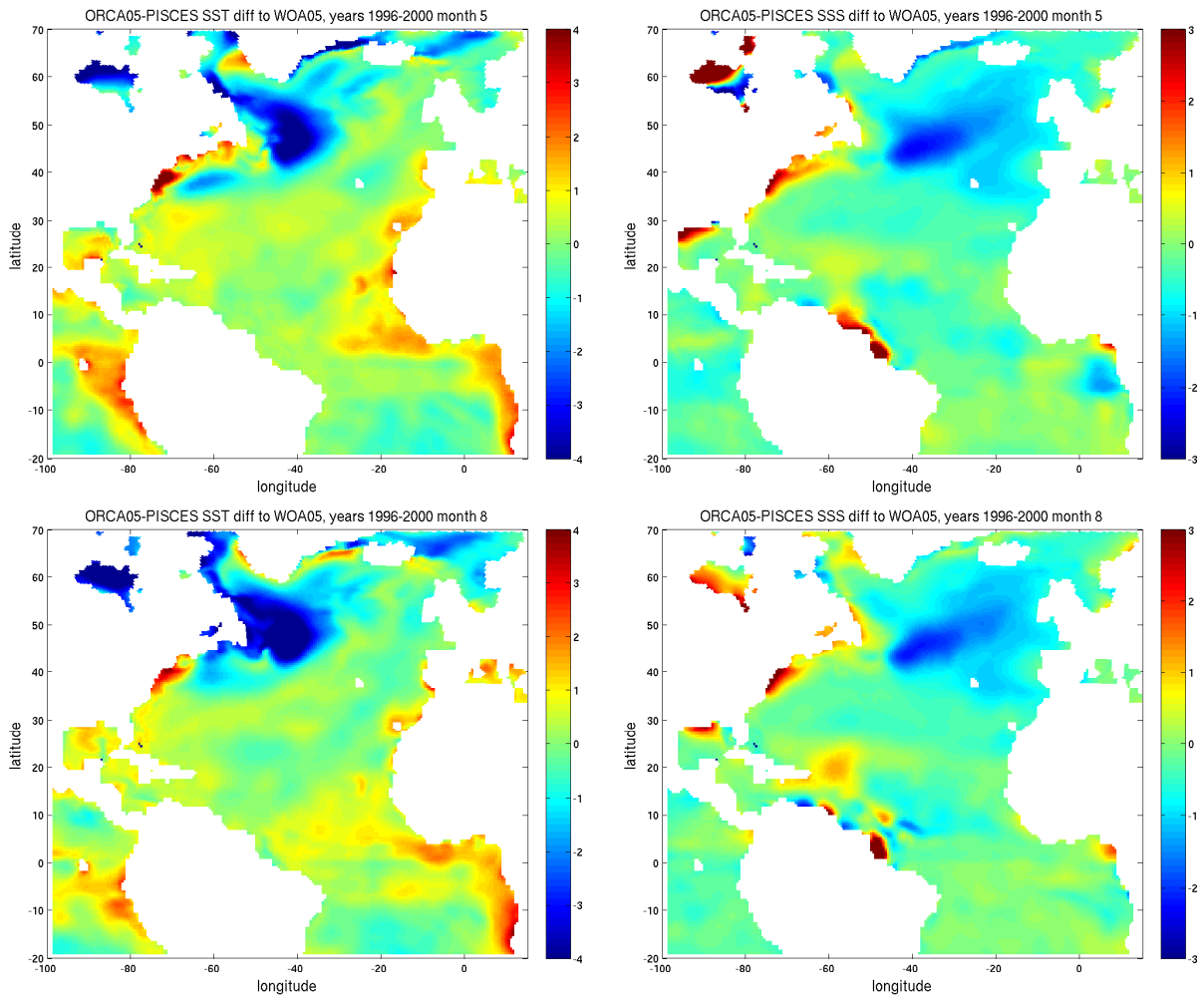


Figure 26: SST & SSS : ORCA05-PISCES diff to WOA05, years 1996-2000 in may and august

Annex 3 : Comparaison of oxygen with World Ocean Atlas

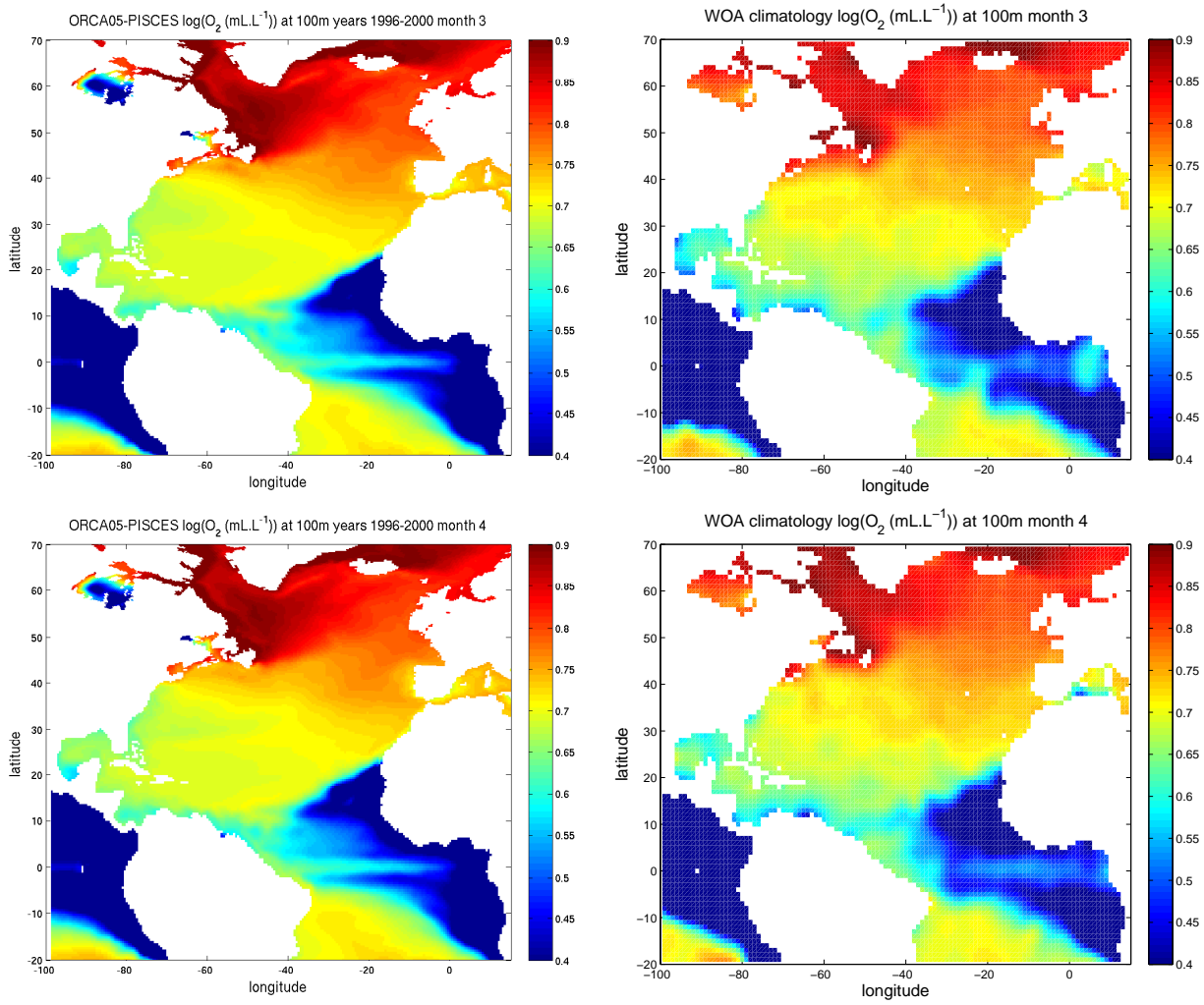


Figure 27: Oxygen at 100m in ORCA05-PISCES years 1996-2000 and WOA5 in march and april

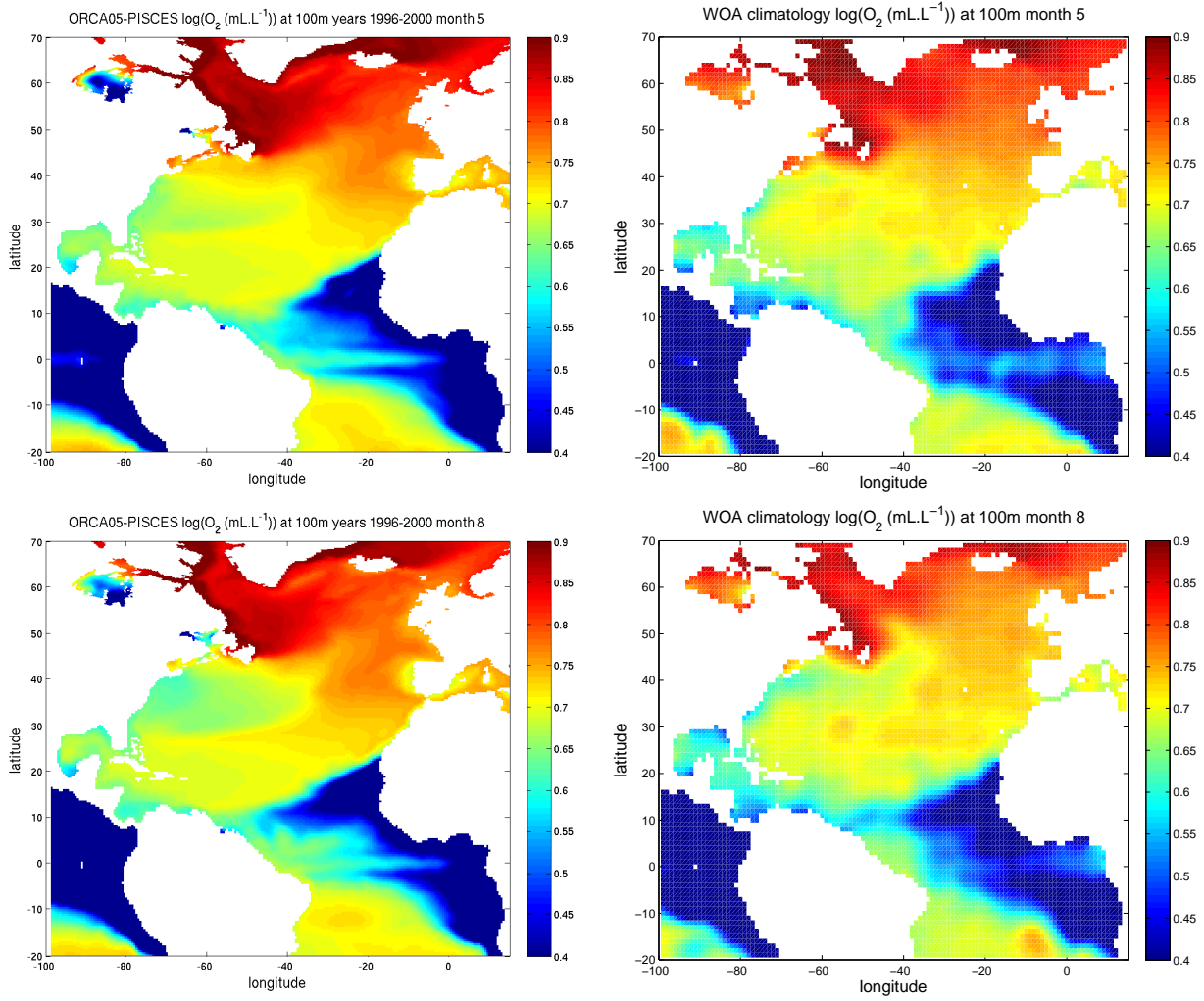


Figure 28: Oxygen at 100m in ORCA05-PISCES years 1996-2000 and WOA05 in may and august

Oxygen sections a 30 °N

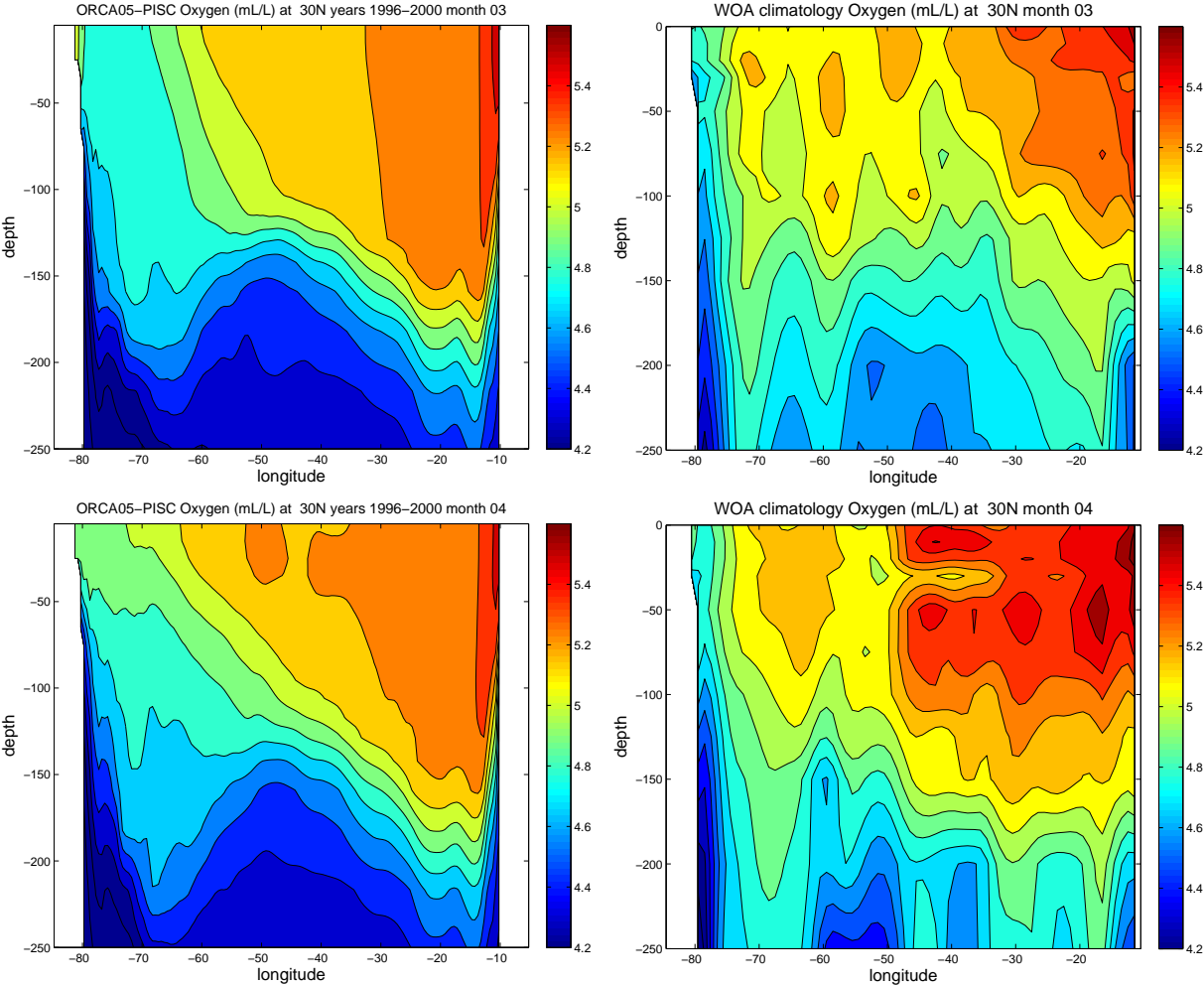


Figure 29: Oxygen section at 30 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in march and april

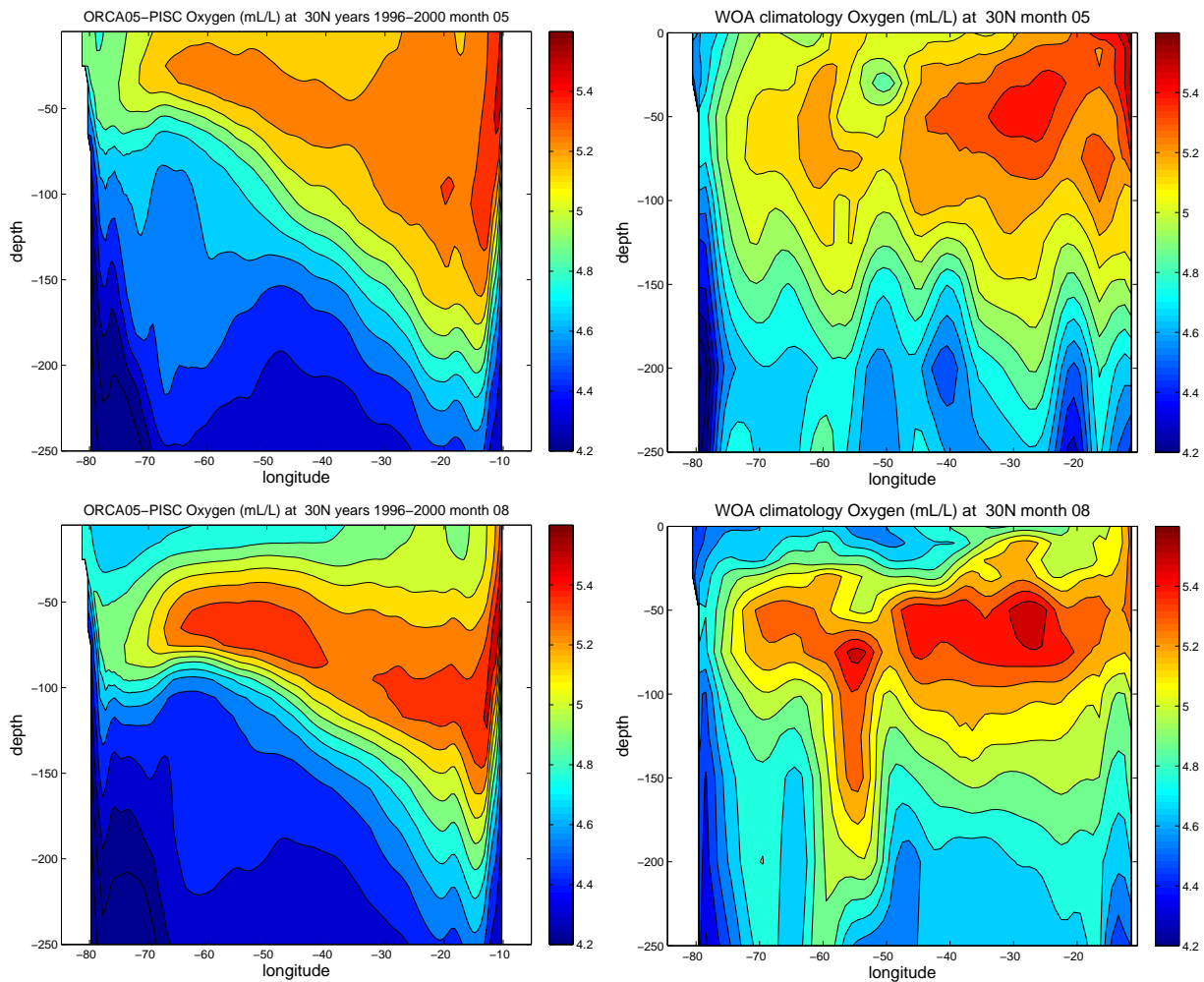


Figure 30: Oxygen section at 30 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in may and august

Oxygen sections at 40 °N

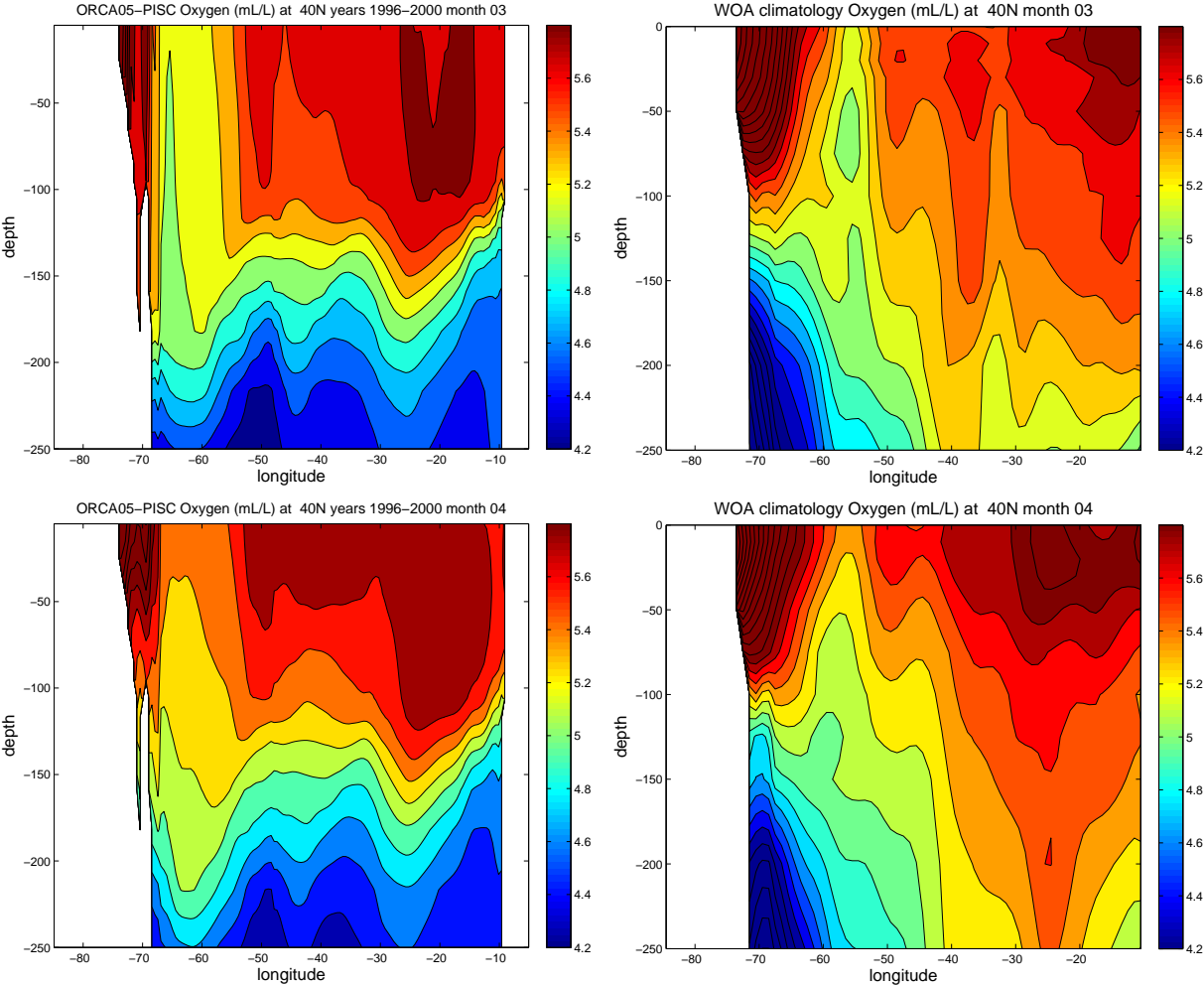


Figure 31: Oxygen section at 40 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in march and april

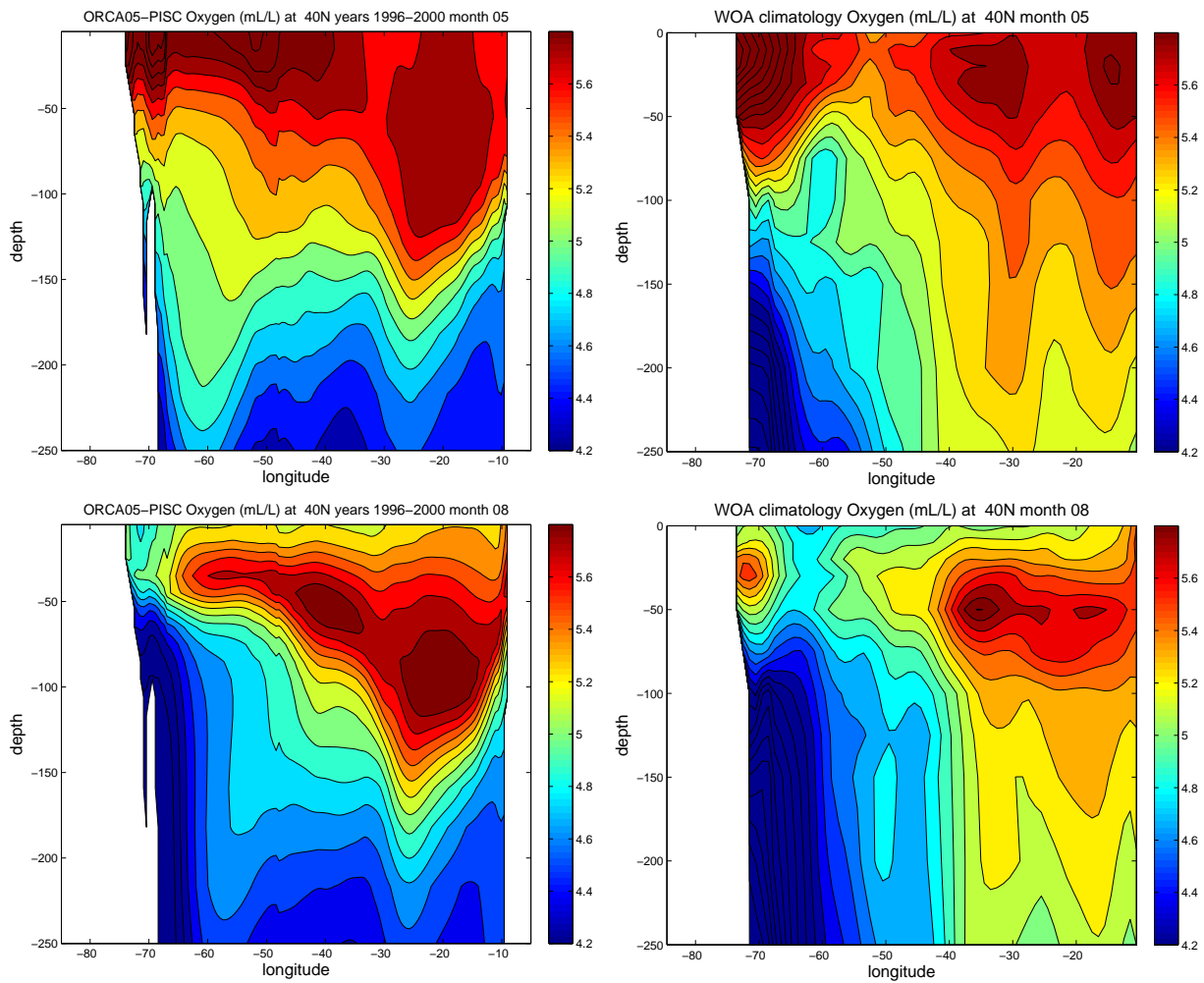


Figure 32: Oxygen section at 40 °N in ORCA05-PISCES years 1996-2000 compared with World Ocean Atlas climatology in may and august

Annex 4 : Chlorophyll plots with linear scale

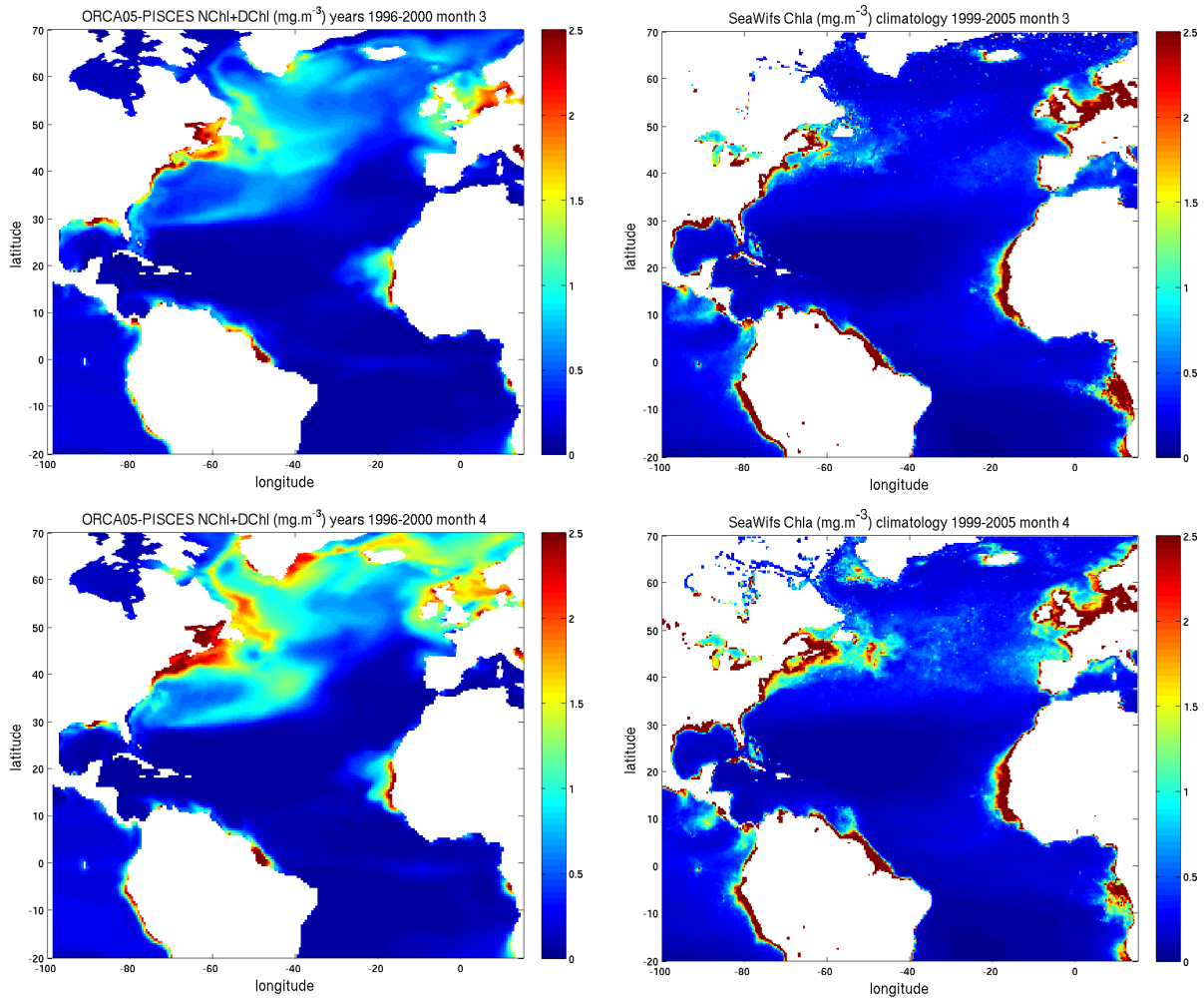


Figure 33: Mean surface chlorophyll in ORCA05-PISCES years 1996-2000 compared with Seawifs climatology in march and april

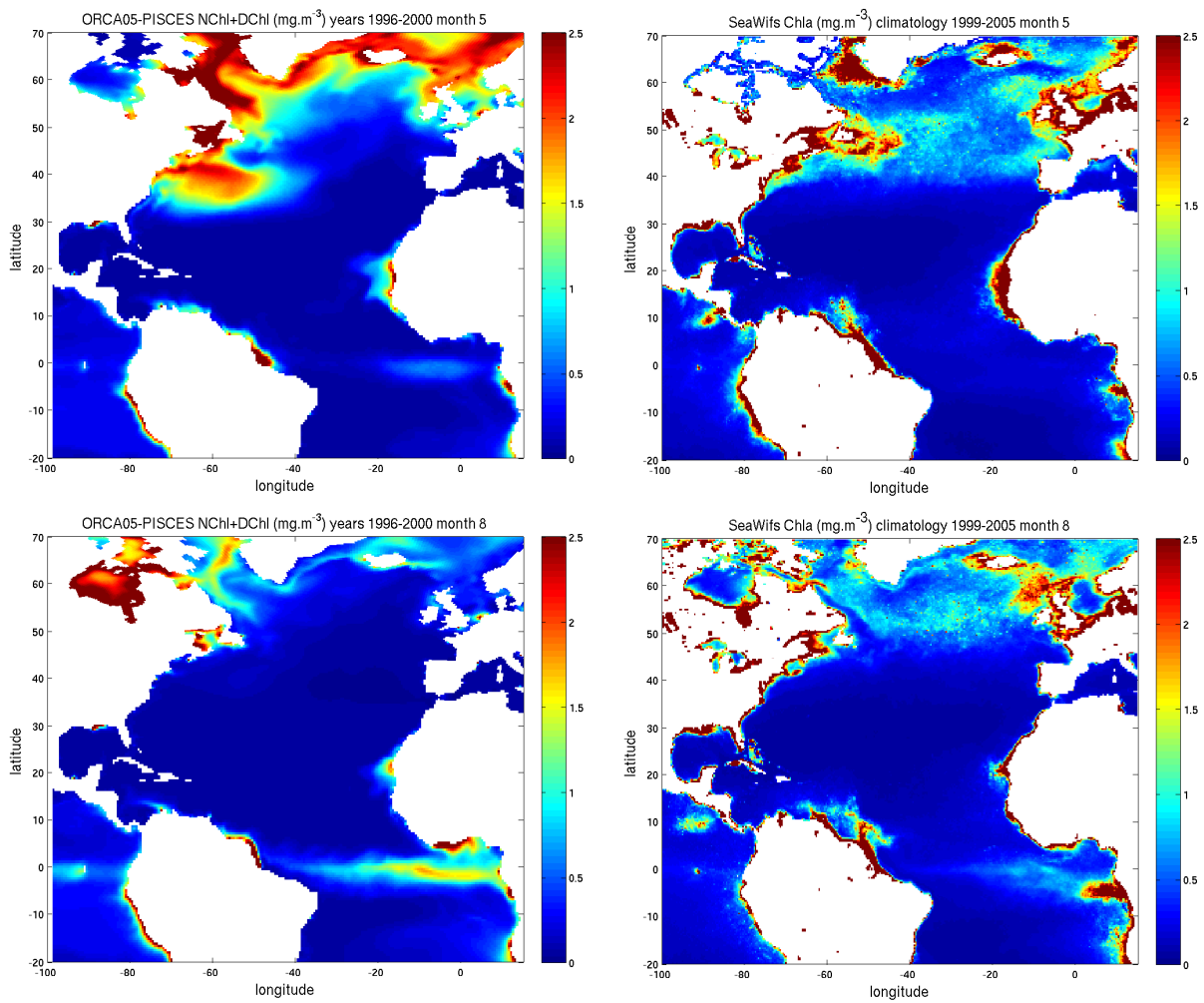


Figure 34: Mean surface chlorophyll in ORCA05-PISCES years 1996-2000 compared with Seawifs climatology in may and august


```

nstock      =    5475
nwrite      =    5475
nrunoff     =      2
ln_ctl      = .false.
nictls     =      0
nictle     =      0
njctls     =      0
njctle     =      0
isplt      =      1
jsplt      =      1
nbench     =      0
/
!-----
!      nam_mpp      Massively Parallel Processing
!-----
!  c_mpi_send      mpi send/recieve type
!                  = 'S' : standard blocking send
!                  = 'B' : buffer blocking send
!                  = 'I' : immediate non-blocking send
&nam_mpp
  c_mpi_send = 'S'
/
!-----
!      nam_traadv   advection scheme for tracer (option not control by CPP keys)
!-----
!  ln_traadv_cen2   2nd order centered scheme      (default T)
!  ln_traadv_tvd    TVD scheme                      (default F)
!  ln_traadv_muscl  MUSCL scheme                   (default F)
!  ln_traadv_muscl2 MUSCL2 scheme                   (default F)
&nam_traadv
  ln_traadv_cen2 = .false.
  ln_traadv_tvd  = .true.
  ln_traadv_muscl = .false.
  ln_traadv_muscl2 = .false.
/
!-----
!      nam_traldf   lateral diffusion scheme for tracer (option not control by CPP keys)
!-----
!  Type of the operator :
!  ln_traldf_lap     laplacian operator             (default T)
!  ln_traldf_bilap   bilaplacian operator          (default F)
!  Direction of action :
!  ln_traldf_level   iso-level                     (default F)
!  ln_traldf_hor     horizontal (geopotential)     (default F)^**
!  ln_traldf_iso     iso-neutral                   (default T)^*
!  Coefficient
!  aht0              horizontal eddy diffusivity for tracers (m2/s)
!  ahtb0             background eddy diffusivity for isopycnal diffusion (m2/s)
!  aeiv0             eddy induced velocity coefficient (m2/s)
!  ^* require key_ldfslp to compute the direction of the lateral diffusion
!  ^** require key_ldfslp in s-coordinate
&nam_traldf

```

```

ln_traldf_lap    = .true.
ln_traldf_bilap = .false.
ln_traldf_level = .false.
ln_traldf_hor   = .false.
ln_traldf_iso   = .true.
aht0    = 1000.
ahtb0   = 0.
aeiv0   = 1000.
/
!-----
!      nam_dynldf   lateral diffusion on momentum
!-----
! Type of the operator :
!   ln_dynldf_lap   laplacian operator           (default T)
!   ln_dynldf_bilap bilaplacian operator        (default F)
! Direction of action :
!   ln_dynldf_level iso-level                   (default F)
!   ln_dynldf_hor   horizontal (geopotential)   (default F)^**
!   ln_dynldf_iso   iso-neutral                 (default T)^*
! Coefficient
!   ahm0    horizontal eddy viscosity for the dynamics (m2/s)
!   ahmb0   background eddy viscosity for isopycnal diffusion (m2/s)
&nam_dynldf
  ln_dynldf_lap    = .false.
  ln_dynldf_bilap = .true.
  ln_dynldf_level = .true.
  ln_dynldf_hor   = .false.
  ln_dynldf_iso   = .false.
  ahm0    = -8.5e+11
  ahmb0   = 0.
/
!-----
!      namflg      algorithm flags (algorithm not control by CPP keys)
!-----
!   ln_dynhpg_imp  hydrostatic pressure gradient: semi-implicit time scheme (T)
!                                     centered      time scheme (F)
&namflg
  ln_dynhpg_imp = .true.
/
!-----
!      nam_dynvor  option of physics/algorithm (not control by CPP keys)
!-----
!   ln_dynvor_ens  vorticity trends: enstrophy conserving scheme (default T)
!   ln_dynvor_ene  "      " : energy conserving scheme (default F)
!   ln_dynvor_mix  "      " : mixed scheme (default F)
!   ln_dynvor_eeen "      " : energy & enstrophy scheme (default F)
&nam_dynvor
  ln_dynvor_ene = .FALSE.
  ln_dynvor_ens = .FALSE.
  ln_dynvor_mix = .FALSE.
  ln_dynvor_eeen = .TRUE.
/

```



```

!-----
!      namtau   surface wind stress
!-----
!  ntau000   gently increase the stress over the first ntau_rst time-steps
!  tau0x     uniform value used as default surface heat flux
!  tau0y     uniform value used as default solar radiation flux
&namtau
  ntau000 =      0
  tau0x   =      0.e0
  tau0y   =      0.e0
/
!-----
!      namflx   surface fluxes
!-----
!  q0        uniform value used as default surface heat flux
!  qsr0      uniform value used as default solar radiation flux
!  emp0      uniform value used as default surface freshwater budget (E-P)
&namflx
  q0       =      0.e0
  qsr0     =      0.e0
  emp0     =      0.e0
/
!-----
!      namalb   albedo parameters
!-----
!  cgren     correction of the snow or ice albedo to take into account
!  albice    albedo of melting ice in the arctic and antarctic
!  alphd     coefficients for linear interpolation used to compute albedo
!            between two extremes values (Pyane, 1972)
!  alphc     "
!  alphdi    "
&namalb
  cgren    =      0.06
  albice   =      0.5
  alphd    =      0.80
  alphc    =      0.65
  alphdi   =      0.72
/
!-----
!      namdom   space and time domain (bathymetry, mesh, timestep)
!-----
!  ntopo     = 0/1 ,compute/read the bathymetry file
!            (mbathy, nb of T-ocean levels)
!  e3zps_min the thickness of the partial step is set larger than the
!  e3zps_rat the minimum of e3zps_min and e3zps_rat * e3t
!            (N.B. 0<e3zps_rat<1)
!  ngrid     = 0/1, compute/read the horizontal mesh
!            (coordinates, scale factors)
!  nmsh      =1 create a mesh file (coordinates, scale factors, masks)
!  nacc      the acceleration of convergence method
!            = 0, no acceleration, rdt = rdttra
!            = 1, acceleration used, rdt < rdttra(k)

```

```

! atfp      asselin time filter parameter
! rdt       time step for the dynamics (and tracer if nacc=0)
! rdtmin    minimum time step on tracers
! rdtmax    maximum time step on tracers
! rdth      depth variation of tracer time step
! nface     frequency of ice model call
! nfbulk    frequency of bulk formulae call (not used if ice used)
! nclosea   = 0 no closed sea
!           = 1 closed sea (Black Sea, Caspian Sea, Great US Lakes...)
&namdom
  ntopo     =      1
  e3zps_min =     25.
  e3zps_rat =     0.2
  ngrid     =      1
  nmsh      =      1
  nacc      =      0
  atfp      =     0.1
  rdt       =    2160.
  rdtmin    =    2160.
  rdtmax    =    2160.
  rdth      =     800.
  nface     =      5
  nfbulk    =      5
  nclosea   =      0
/
!-----
!          namfwb  freshwater budget correction
!-----
! ln_fwb     logical flag for freshwater budget correction (0 annual mean)
&namfwb
  ln_fwb     = .false.
/
!-----
!          namptr  Poleward Transport Diagnostic
!-----
! ln_diaptr  logical flag for Poleward transport computation
! nf_ptr     Frequency of computation
&namptr
  ln_diaptr = .false.
  nf_ptr    = 15
/
!-----
!          namcro  cross land advection
!-----
! n_cla     advection between 2 ocean pts separates by land
&namcla
  n_cla     = 0
/
!-----
!          namzdf  vertical physics
!-----
! ln_zdfevd enhanced vertical diffusion          (default T)

```

```

! ln_zdfnpc Non-Penetrative Convection          (default T)
! avm0      vertical eddy viscosity for the dynamic (m2/s)
! avt0      vertical eddy diffusivity for tracers (m2/s)
! avevd     vertical coefficient for enhanced diffusion scheme (m2/s)
! nevdm     = 0  apply enhanced mixing on tracer only
!           = 1  apply enhanced mixing on both tracer and momentum
! ln_zdfexp vertical physics: (=T) time splitting (T)      (Default=F)
!           (=F) euler backward (F)
! n_zdfexp  number of sub-timestep for time splitting scheme
&namzdf
  ln_zdfevd = .true.
  ln_zdfnpc = .false.
  avm0      = 1.e-4
  avt0      = 1.e-5
  avevd     = 10.
  nevdm     = 1
  ln_zdfexp = .false.
  n_zdfexp  = 3
/
!-----
!          namnpc  vnon penetrative convection
!-----
! nnpc1     non penetrative convective scheme frequency
! nnpc2     non penetrative convective scheme print frequency
&namnpc
  nnpc1 = 1
  nnpc2 = 365
/
!-----
!          namdbl  bottom boundary layer scheme
!-----
! atrdbl    lateral tracer coeff. for bottom boundary layer scheme(m2/s)
&namdbl
  atrdbl = 6000.
/
!-----
!          namric  richardson number dependent vertical diffusion
!                 ( #ifdef "key_zdfrichardson" )
!-----
! avmri     maximum value of the vertical viscosity
! alp       coefficient of the parameterization
! nric      coefficient of the parameterization
&namric
  avmri = 100.e-4
  alp   = 5.
  nric  = 2
/
!-----
!          namtke  turbulent eddy kinetic dependent vertical diffusion
!                 ( #ifdef "key_zdftke" )
!-----
! ln_rstke  flag to restart with tke from a run without tke (default F)

```

```

! ediff    coef. to compute vertical eddy coef. (avt=ediff*mxl*sqrt(e) )
! ediss    coef. of the Kolmogoroff dissipation
! ebb      coef. of the surface input of tke
! efave    coef. to applied to the tke diffusion ( avtke=efave*avm )
! emin     minimum value of tke (m^2/s^2)
! emin0    surface minimum value of tke (m^2/s^2)
! nitke    number of restart iterative loops
! ri_c     critic richardson number
! nmxl     flag on mixing length used
!          = 0 bounded by the distance to surface and bottom
!          = 1 bounded by the local vertical scale factor
!          = 2 first vertical derivative of mixing length bounded by 1
! npdl     flag on prandtl number
!          = 0 no vertical prandtl number (avt=avm)
!          = 1 prandtl number function of richarson number (avt=pdl*avm)
!          = 2 same as = 1 but a shapiro filter is applied on pdl
! nave     = horizontal averaged (=1) or not (=0) of avt (default =1)
! navb     = 0 cst background avt0, avm0 / =1 profile used on avtb

```

```
&namtke
```

```

ln_rstke = .false.
ediff =    0.1
ediss =    0.7
ebb =      60
efave =    1.
emin =    1.e-6
emin0 =    1.e-4
nitke =    50
nmxl =     2
npdl =     1
navb =     0
nave =     1
ln_lsfc = .true.
lmin  =    0.4
n_etau =    1
nhtau =    1
fr_emin =  0.05
ln_lc = .true.
c_lc =    0.15

```

```
/
```

```

!-----
!          namkpp    K-Profile Parameterization dependent vertical diffusion
!                   ( #ifdef "key_zdfkpp" )
!-----

```

```

! ln_kpprimix    shear instability mixing (default T)
! difmiw         constant internal wave viscosity (m2/s)
! difsiw         constant internal wave diffusivity (m2/s)
! Riinfy         local Richardson Number limit for shear instability
! difri          maximum shear mixing at Rig = 0 (m2/s)
! bvsqcon        Brunt-Vaisala squared (1/s**2) for maximum convection
! difcon         maximum mixing in interior convection (m2/s)
! nave          = 0/1 flag for horizontal average on avt, avmu, avmv
! navb          = 0/1 flag for constant or profile background avt

```

```

&namkpp
  ln_kpprimix = .true.
  difmiw      = 1.e-04
  difsiw      = 0.1e-04
  Riinfy      = 0.8
  difri       = 0.0050
  bvsqcon     = -0.01e-07
  difcon      = 1.
  navb        = 0
  nave        = 1
/
!-----
!      namddm  double diffusive mixing parameterization
!-----
!  avts      maximum avs for dd mixing
!  hsbfr     heat/salt buoyancy flux ratio
&namddm
  avts = 1.e-4
  hsbfr = 1.6
/
!-----
!      namlbc  lateral momentum boundary condition
!-----
!  shlat     lateral boundary condition on velocity
!            shlat = 0 , free slip
!            0 < shlat < 2 , partial slip
!            shlat = 2 , no slip
!            2 < shlat      , strong slip
&namlbc
  shlat = 0.
/
!-----
!      nambfr  bottom friction
!-----
!  nbotfr    type of bottom friction
!            nbotfr = 0 , no slip
!            nbotfr = 1 , linear friction
!            nbotfr = 2 , nonlinear friction
!            nbotfr = 3 , free slip
!  bfri1     bottom drag coefficient (linear case)
!  bfri2     bottom drag coefficient (non linear case)
!  bfeb2     bottom turbulent kinetic energy (m^2/s^2)
&nambfr
  nbotfr = 2
  bfri1 = 4.e-4
  bfri2 = 1.e-3
  bfeb2 = 2.5e-3
/
!-----
!      nambbc  bottom temperature boundary condition
!-----
!  ngeo_flux = 0 no geothermal heat flux

```

```

!           = 1 constant geothermal heat flux
!           = 2 variable geothermal heat flux (read in geothermal_heating.nc)
!           ( C A U T I O N : flux in mW/m2 in the NetCDF file )
! ngeo_flux_const  Constant value of geothermal heat flux (W/m2)
&nambbc
  ngeo_flux = 1
  ngeo_flux_const = 88.4e-3
/
!-----
!           namqsr  penetrative solar radiation
!-----
! ln_traqsr : penetrative solar radiation (T) or not (F)      (Default=T)
! rabs      fraction of qsr associated with xsi1
! xsi1     first depth of extinction
! xsi2     second depth of extinction
&namqsr
  ln_traqsr = .true.
  rabs      = 0.58
  xsi1     = 0.35
  xsi2     = 23.0
  ln_qsr_sms = .false.
/
!-----
!           namtdp  tracer newtonian damping ('key_tradmp')
!-----
! ndmp     type of damping in temperature and salinity
!           (= 'latitude', damping poleward of 'ndmp' degrees and function
!           of the distance-to-coast. Red and Med Seas as ndmp=-1)
!           (= -1 damping only in Med and Red Seas)
! ndmpf    =1 create a damping.coeff NetCDF file (the 3D damping array)
! nmldmp   type of damping in the mixed layer
!           (=0 damping throughout the water column)
!           (=1 no damping in the mixed layer defined by avt >5cm2/s )
!           (=2 no damping in the mixed layer defined rho<rho(surf)+.01 )
! sdmp     surface time scale for internal damping (days)
! bdmp     bottom time scale for internal damping (days)
! hdmp     depth of transition between sdmp and bdmp (meters)
&namtdp
  ndmp     = -1
  ndmpf    = 1
  nmldmp   = 0
  sdmp     = 50.
  bdmp     = 360.
  hdmp     = 800.
/
!-----
!           nameos  ocean physical parameters
!-----
! neos     type of equation of state and Brunt-Vaisala frequency
!           = 0, UNESCO (formulation of Jackett and McDougall (1994)
!           and of McDougall (1987) )
!           = 1, linear: rho(T) = rau0 * ( 1.028 - ralpha * T )

```

```

!           = 2, linear: rho(T,S) = rau0 * ( rbeta * S - ralpha * T )
!
!           with rau0=1020 set in parcst routine
!  ralpha  thermal expansion coefficient (linear equation of state)
!  rbeta   saline expansion coefficient (linear equation of state)
&nameos
  neos    =      0
  ralpha  =  2.e-4
  rbeta   =  0.001
/
!-----
!           namsol  elliptic solver / island / free surface
!-----
!  nsolv   elliptic solver (=1 preconditioned conjugate gradient: pcg)
!           (=2 successive-over-relaxation: sor)
!           (=3 FETI: fet, all require "key_feti" defined)
!           (=4 sor with extra outer halo))
!  nsol_arp absolute/relative (0/1) precision convergence test
!  nmin    minimum of iterations for the SOR solver
!  nmax    maximum of iterations for the SOR solver
!  nmod    frequency of test for the SOR solver
!  eps     absolute precision of the solver
!  resmax  absolute precision for the SOR solver
!  sor     optimal coefficient for SOR solver
!  epsisl  absolute precision on stream function solver
!  nmisl   maximum pcg iterations for island
!  rnu     strength of the additional force used in free surface b.c.
&namsol
  nsolv    =      2
  nsol_arp =      0
  nmin     =      0
  nmax     =     800
  nmod     =      5
  eps      =  1.E-6
  resmax   =  1.E-11
  sor      =  1.964
  epsisl   =  1.e-10
  nmisl    =  4000
  rnu      =      1.
/
!=====
!  Diagnostics namelists
!  namtrd  dynamics and/or tracer trends
!  namgap  level mean model-data gap
!  namznl  zonal mean heat & freshwater fluxes computation
!  namspr  surface pressure in rigid-lid
!=====
!-----
!  namtrd  diagnostics on dynamics and/or tracer trends
!           ('key_diatrdyn' and/or 'key_diatrtra')
!           or mixed-layer trends ('key_diatrdmld')
!-----
!  ntrd    time step frequency dynamics and tracers trends

```



```

! nctls control surface type in mixed-layer trends (0,1 or n<jpk)
&namtrd
  ntrd = 365
  nctls = 0
/
!-----
!          namgap    level mean model-data gap ('key_diagap')
!-----
! ngap    time-step frequency of model-data gap computation
! nprg    time-step frequency of gap print in model output
&namgap
  ngap = 15
  nprg = 10
/
!-----
!          namznl    zonal mean heat & freshwater fluxes computation
!                   (#ifdef "key_diaznl")
!-----
! nfznl    time-step frequency of zonal mean fluxes computation
&namznl
  nfznl = 15
/
!-----
!          namspr    surface pressure diagnostic
!-----
! nmaxp    maximum of iterations for the solver
! epsp     absolute precision of the solver
! nitersp  number of iteration done by the solver
&namspr
  nmaxp = 1000
  epsp  = 1.e-3
  nitersp = 400
/
!-----
!          namcpl    coupled ocean/atmosphere model (#ifdef "key_coupled")
!-----
! nexco    coupling frequency in time steps
! cchan    coupling technique 'PIPE' or 'CLIM'
&namcpl
  nexco      = 24
  cchan      = 'PIPE'
  nmodcpl    = 2
  cplmodnam  = 'opa.xx'
  cploasis   = 'Oasis'
  nfldo2c    = 2
  nflxc2o    = 6
  ntauc2o    = 4
  cpl_writ(1) = 'SOSSTSST'
  cpl_f_writ(1) = 'ocesst'
  cpl_writ(2) = 'SOICECOV'
  cpl_f_writ(2) = 'oceice'
  cpl_readflx(1) = 'SONSFLDO'

```

```

cpl_f_readflx(1) = 'oceflx'
cpl_readflx(2)  = 'SOSHFLDO'
cpl_f_readflx(2) = 'oceflx'
cpl_readflx(3)  = 'SOTOPRSU'
cpl_f_readflx(3) = 'oceflx'
cpl_readflx(4)  = 'SOTFHSU'
cpl_f_readflx(4) = 'oceflx'
cpl_readflx(5)  = 'SORUNCOA'
cpl_f_readflx(5) = 'oceflx'
cpl_readflx(6)  = 'SORIVFLU'
cpl_f_readflx(6) = 'oceflx'
cpl_readtau(1)  = 'SOZOTAUX'
cpl_f_readtau(1) = 'ocetau'
cpl_readtau(2)  = 'SOZOTAU2'
cpl_f_readtau(2) = 'ocetau'
cpl_readtau(3)  = 'SOMETAU1'
cpl_f_readtau(3) = 'ocetau'
cpl_readtau(4)  = 'SOMETAU2'
cpl_f_readtau(4) = 'ocetau'
/
!-----
!      namobc      open boundaries parameters (#ifdef key_obc)
!-----
! nobc_dta   = 0 the obc data are equal to the initial state
!             = 1 the obc data are read in 'obc .dta' files
! rdpeob    time relaxation (days) for the east open boundary
! rdpwob    time relaxation (days) for the west open boundary
! rdpnob    time relaxation (days) for the north open boundary
! rdpsob    time relaxation (days) for the south open boundary
! zbsic1    barotropic stream function on isolated coastline 1
! zbsic2    barotropic stream function on isolated coastline 2
! zbsic3    barotropic stream function on isolated coastline 3
&namobc
  nobc_dta = 0
  rdpein   = 1.
  rdpwin   = 1.
  rdpnin   = 30.
  rdpsin   = 1.
  rdpeob   = 1500.
  rdpwob   = 15.
  rdpnob   = 150.
  rdpsob   = 15.
  zbsic1   = 140.e+6
  zbsic2   = 1.e+6
  zbsic3   = 0.
/
!-----
!      namflo      float parameters (#ifdef key_float)
!-----
! ln_rstflo  boolean term for float restart (true or false)
! nwritefl   frequency of float output file
! nstockfl   frequency of float restart file

```

```

! ln_argo      Argo type floats (stay at the surface each 10 days)
! ln_flork4   = T trajectories computed with a 4th order Runge-Kutta
!             = F (default)   computed with Blanke' scheme
&namflo
  ln_rstflo = .false.
  nwritefl  =      75
  nstockfl  =    5475
  ln_argo   = .false.
  ln_flork4 = .false.
/

```

Annex 6 : CPP Keys

```

key_mpp_mpi
key_partial_steps
key_trabbl_dif
key_vectopt_loop
key_vectopt_memory
key_orca_r05
key_ice_lim
key_lim_fdd
key_dynspg_fsc
key_diaeiv
key_ldfslp
key_traldf_c2d
key_traldf_eiv
key_dynldf_c2d
key_dtatemp
key_dtasal
key_tau_daily
key_flx_bulk_daily
key_tradmp
key_trabbc
key_zdftke
key_zdfddm
key_passivetr
key_trc_pisces
key_trcbbl_dif
key_trc_zdfddm
key_trc_dta
key_trc_ldfeiv
key_trc_dia3d
key_trc_diaadd

```